

# GROUNDWATER RESOURCES OF LOUDOUN COUNTY, VIRGINIA

by

J. Roy Murphy

NORTHERN REGIONAL OFFICE



COMMONWEALTH OF VIRGINIA

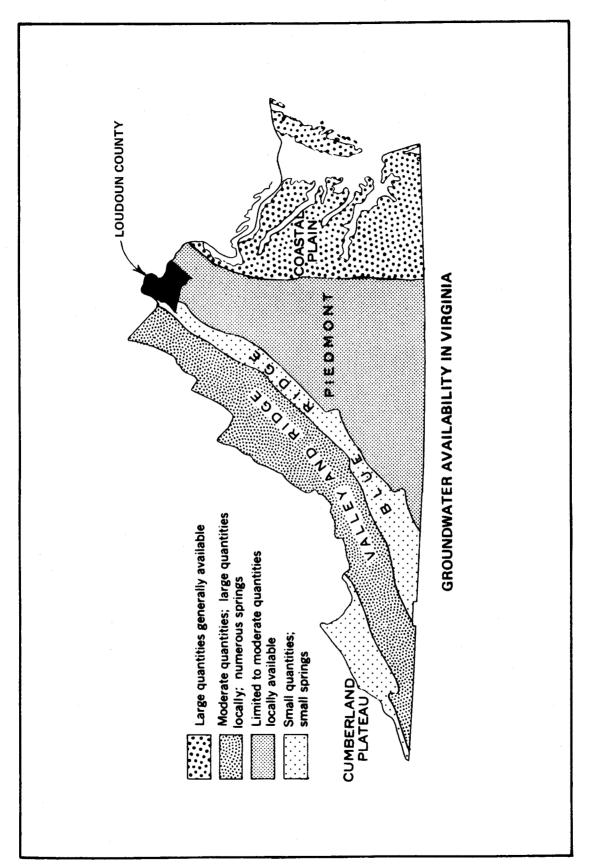
STATE WATER CONTROL BOARD

BUREAU OF WATER CONTROL MANAGEMENT

Richmond, Virginia

Planning Bulletin 315

**July 1979** 



# LOUDOUN COUNTY GROUNDWATER PRESENT CONDITIONS AND PROSPECTS



by J. R. Murphy Northern Regional Office

VIRGINIA STATE WATER CONTROL BOARD BUREAU OF WATER CONTROL MANAGEMENT

PLANNING BULLETIN 315

JULY 1979

		•
		•
		, k

# TABLE OF CONTENTS

	Page
LIST OF PLATES	
LIST OF TABLES	
LIST OF FIGURES	
ACKNOWLEDGEMENTS	х
FORWARD	xii
ABSTRACT	xiv
CHAPTER	
I INTRODUCTION	1
Background	1
Purpose and Scope	5
Methods of Investigation	5
Previous Investigations	6
Water Well Numbering Syst	tem 7
II PHYSICAL SETTING	9
Physiography	9
Hydrology	10
Climate	17
Soils	21
Vegetation	26
III HYDROGEOLOGY	27
Geologic Setting	27
Geologic Formations	33
Geologic Structure	42
Groundwater	43

# TABLE OF CONTENTS

		Page
СНАРТЕ	CR CR	
III	Aquifers	48
	Basement Complex - Blue Ridge Province	50
	Layered Clastics	51
	Layered Volcanics	52
	Triassic Lowlands	53
IV	GROUNDWATER QUALITY	55
	Constituents in Natural Groundwater	55
	Water Quality in Loudoun County	62
	Triassic Lowlands	63
	Layered Clastics	63
	Layered Volcanics	64
	Basement Complex	64
	Groundwater Contamination	67
٧	GROUNDWATER DEVELOPMENT	73
	General Setting	73
	Public Supplies	74
	Key Water Wells	75
	The Groundwater Act of 1973	76
	Groundwater Availability	79
	Well Depths	8.4

# TABLE OF CONTENTS

		Page
СНАРТЕ	R	
VI	CONCLUSIONS AND RECOMMENDATIONS	87
	APPENDIX A	A-1
٠	Summary of Groundwater Quality Analyses for	A-3
	Loudoun County	
	APPENDIX B	B-1
	Summary of Water Well Data for Loudoun County	B-3
	APPENDIX C	C-1
	Using Large Diameters to Increase Well Storage	C-1
	GLOSSARY OF TERMS	
	RIRI TOCDADHY	

				-
				•
				•
				*
				ζ.

# LIST OF PLATES

Plate No.		Page
	Groundwater Availability in Virginia	Frontispiece
1	Location Map of Loudoun County	3
2	Physiography of Loudoun County	11
3	Major Watershed of Loudoun County	13
4	Soil Groupings	19
5	Generalized Geologic Map of Loudoun County	29
6	Hydrologic Cycle and Aquifers	45
7	Trends of Groundwater Hardness in Loudoun	65
	County	
8	Key Water Wells in Loudoun County	77
9	Groundwater Availability in Loudoun County	81
	Virginia State Water Control Board Offices	Fnd Flan

# LIST OF TABLES

Table	No.	Page
1	Average Discharge at Main River Gaging Stations	16
2	Soil Groupings	24
3	Geologic Formations and Their Water-Bearings	34
	Properties	
4	Virginia Groundwater Standards (for human	56
	consumption)	
5	Summary of Analyses of Groundwater	71
6	Well Depth Summary	86

		•
		•
		•
		*
		~
		*.

# LIST OF FIGURES

Figure No.		Page
1	Idealized Soil Profile in a Region of	22
	Mild Temperature and Moderate Rainfall	
2	Soil Moisture Distribution	44

#### **ACKNOWLEDGEMENTS**

Appreciation is expressed to the citizens of Loudoun County who provided information to employees of the State Water Control Board on wells which they own and who allowed water samples to be taken. Also, officials of Leesburg, the County Health Department, AT&T and other agencies which supplied information are to be thanked. Particular thanks goes to those water well contractors who worked in Loudoun County and faithfully have rendered water well completion reports. Lastly the summer employees of the State Water Control Board, particularly Haywood A. Wigglesworth, who though temporary, rendered dedicated and vigilant service in sample collection, analysis, and data collection are to be praised.

			*
			•
			•
			•
			·
			<b>*</b> .
			*
			4

#### **FOREWARD**

This preliminary report is part of a series intended ultimately to cover the Commonwealth. The purpose is to provide private citizens, groundwater users, developers, well drilling contractors, consultants, governments and other users with as complete a summary, as possible, of the groundwater situation, within the geologic constraints, in each county.

Using this report as a basic document, prospective groundwater users and other interested parties may form estimates of the potential occurrence of subsurface water. Any detailed studies or point-source studies should, of course, be performed by a trained geohydrologist. The State Water Control Board will be available as a public service organization to accumulate future data, act as source of general information and to perform a future update of the groundwater resources and quality. A glossary of terms is provided to aid those persons not familiar with the vocabulary of the geologic profession.

# LOUDOUN COUNTY GROUNDWATER PRESENT CONDITIONS AND PROSPECTS

by

J. R. Murphy

#### **ABSTRACT**

Groundwater is the end-product of geohydrologic conditions and depends on several variables. The greatest variable is the geology, followed closely by the topographic position. In Loudoun County the geologic units range in age from approximately 1.1 billion-year-old granites to 200 million-year-old lava materials.

The water-holding and subsequent yielding characteristics of a geologic unit are a result of the lithology and degree of tectonic deformation. These characteristics determine whether a rock unit is an aquifer (water-bearer) or aquitard (water-barrier) or aquifuge (unproductive). The areas of limestone conglomerate, sandstone, and other sedimentary rocks of the Triassic Basin are all fair-to-excellent producing units. The recent igneous rocks (diabase, etc.), on the other hand, tend to be poor producers, if at all. The areas containing the ancient plutonic and volcanic rocks yield the greatest quantities of water when fault structures are encountered; otherwise they tend to be dry. Other characteristics relevant to a unit's water-bearing potential are discussed in the text.

Data are presented in the text discussing hard and soft water regimes and specific chemical characteristics encountered in each groundwater province.

Pollution must be considered as a threat to continued use of groundwater. Such items as septic tanks, drainfields, buried fuel and chemical tanks, all have a potential for failure. Unlike surface waters, there is no easy, or inexpensive, way to flush the water table. Pollution, regardless of type, normally will stay until it is no longer dangerous. This pollution may last generations, not just years.

The U. S. Census Service estimates that there are approximately 5,000 wells in Loudoun County. Of this number, many are owned and have been used by towns, subdivisions, schools and industries. The majority, however, belong to individuals.

A map is included in this report that indicates the potential for groundwater as it presently is known. Indications are that the major fault zones, the Triassic sediments and the Potomac River margin possess the greatest potential for future groundwater development. Attention also is directed towards several areas wherein larger yields of water have been found than the regional averages would cause one to expect.

#### CHAPTER I

#### INTRODUCTION

#### Background

Loudoun County encompasses 1339 square kilometers (517 square miles) in area, located in the northern tip of Virginia (Plate 1). It occupies a part of economic region III as defined by the Virginia Division of Planning. The eastern edge of the county lies approximately 35 miles west of the nation's capital. The county is rural in nature, although about 35 percent of the population lives in urban areas. One-fourth of the people live in the eastern portion of the county, and slightly less than 50 percent of the employed workers commute outside of the county. The 1975 population was 48,900, with a 5.6 percent increase noted between 1970-74. The incorporated towns and their number of inhabitants (1974 data) are: Hillsboro (98), Leesburg (7000), Lovettsville (485), Middleburg (800), Purcellville (2000), and Round Hill (600). In addition, the planned communities of Sterling Park and Sugarland Run have, respectively, 13,600 and 5,500 persons. The population is expected to increase at a slower rate than previously, to 60,100 in 1980, and ultimately to 110,200 by the year 2000.

Presently about two-thirds of the total acreage of Loudoun County is in agricultural or open space useage. The number of large farms has continued to increase yearly, whereas the number of acres actually used for agriculture has declined. Statewide the county rated ninth (1970) in total value of all agricultural products, with 87 percent being livestock and related products. As a result of local government's

efforts, several industries have moved into the southeast portion of the county around the Dulles Airport Complex. These industries include manufacturing and training. Other labor categories include, in order of the size of the labor force, Public Administration, Retail and Wholesale trade and Services, with smaller, but equal numbers in non-agricultural wage and salary, and Non-manufacturing (Data Summary, Loudoun County, 1975). Professionals make up almost 27 percent of the total labor force, with clerical specialties being second.

The increase in population, plus the change in the economic base, has produced a resultant change in housing patterns. Previously, prior to 1950, single dwellings on one-acre plots, or large plots with scattered apartments, were normal. The practices of today: to develop large apartment complexes and townhouses, plus subdivisions, have produced increased local demands on water-supply facilities.

In eastern Loudoun County most of the water is purchased from the City of Fairfax (Goose Creek Water System). This causes a problem. As the need for water increases, so does the price and the price of procurement. Groundwater already is a partial answer to this problem. The town of Hamilton obtains water from six wells. Leesburg utilizies six wells and a spring, whereas Middleburg only uses groundwater as a supplement to river supplies. Hillsboro, Purcellville, and Round Hill use springs and/or dam-reservoir combinations. The County Sanitation Authority is attempting to provide water to the unincorporated areas, although private sources also are used. Local industries frequently have drilled their own wells.

4 2	۔ نو	•	¥
	-		

Surface water, spring, and dam combinations are usually better long-range solutions to water shortages, but they do cost more than groundwater systems and once exhausted cannot cheaply be replaced.

Purpose and Scope

The purpose of this report is to summarize available data concerning groundwater resources in Loudoun County and provide readers with a concise, though preliminary, booklet of the existing and potential groundwater conditions.

It is a harsh fact that man cannot live without potable water supplies. His life expectancy without it is limited to about three days, depending on environment and degree of exertion. Therefore, this report attempts to identify areas of undeveloped groundwater and present the quality of known supplies. Some problem areas will be outlined.

# Methods of Investigation

It has been said by an unknown philosopher that, "All civilization exists at the indulgence of geology." The occurrence, availability and quality of groundwater are geology-dependent. Therefore, the geologic framework of the County was investigated from available reports and unpublished data. Information on the productivity of water wells was examined and consolidated with the geologic data to establish the correlation between geology and well yields. The influence of topography on well yields also was considered, as were the depths and construction characteristics of the wells. Chemical-quality samples were collected at selected sites to establish the effects of geology, and of man, on groundwater quality. These data

were supplemented with chemical analyses from public water-supply wells, provided by the State Department of Health. Other sources of information on the hydrogeology of the county include data from water-level observation wells, water level measurements from private and public wells, pumpage records, pumping tests and other information provided by well drillers.

All water well information and records of water quality analyses used in this report are in the files of the State Water Control Board headquarters office in Richmond and the Northern Regional Office in Alexandria. These data have been computerized for storage and retrieval and were used to compile Appendices A and B.

# Previous Investigations

The earliest mention of the geological conditions of Loudoun County is found in a metalurgical survey by Hotchkiss, published in 1884. During the next few years, several papers and classical studies were made, among these Rogers, (1884); Darton (1891); Geiger and Keith (1891); Keith (1893, 1894); Darton and Keith (1901) and Williams, Keith and Darton (1901). In 1923, two papers prepared by 0. E. Meinzer began to relate to the groundwater situations. Many other geologic papers, too many to enumerate, were published before R. C. Cady released his <a href="Preliminary Report on Ground-water Resources of Northern Virginia">Preliminary Report on Ground-water Resources of Northern Virginia</a> in 1933. This was followed by the full report in 1938. In the 1950's a number of geology and related groundwater reports followed, written by Werner, Whitaker, and the Bloomers. P. M. Johnston published five reports on the geology and groundwater of the Washington area in the 1960's, which added further

information applicable to the County. Other authors, Mack (1965), Toewe (1966, 1968) and Parker (1968), among them, wrote geology reports on portions of the county during the same period. The 1970's have seen the advent of geophysical data and much more detailed, localized reporting. The most recent report on a portion of the county is one by K. Y. Lee of the U. S. Geological Survey (1977). No one map or report, as yet, covers in detail all of Loudoun County. The State Geological Map prepared by the Division of Mineral Resources (1963) does present the entire county, but at a scale of 1:500,000. This map does not give sufficient detail for truly accurate groundwater studies. The map presented as Plate 5 is the product of combining many sources and is general in nature.

### Water Well Numbering System

The water wells studied for this report are identified by the State Water Control Board's Bureau of Water Control Management well numbers. Each well number consists of two parts, such as 153-121. The first number, 153, denotes the County in which the well is located; in this case Loudoun County. The second number, 121, is a sequential number that refers to a specific well in Loudoun County.

		*
		•
		,
		,
		•

#### CHAPTER II

#### PHYSICAL SETTING

# Physiography

Loudoun County lies in two physiographic provinces, with the percentage in each being debatable. The western portion is part of the Blue Ridge Province, whereas the eastern part is in the Piedmont. The boundary was considered by Toewe, (1966) Fenneman, (1938) and Gathright (personal communication) as being on the fault line of the Bull Run Fault. This is identified as a line running north-northeast from near Aldie to a point near U. S. Highway 15 at the Potomac River. West of the fault line are found Precambrian and Cambrian-aged, crystalline rocks, layered volcanics and layered clastics. One thin unit of Ordovician limestone is also found west of the fault. East of the line exist Triassic-age rocks. Several other authors, among them Dietrich, (1970) and Espenshade, (1970) place the boundary in Loudoun County about two miles east of the West Virginia state border at the edge of the layered volcanic rocks near Bluemont. The writer of this publication prefers the placement of the line at the Bull Run Fault, with a further subdivision of the Blue Ridge Province into a Layered Clastics zone, a Layered Volcanics zone, and a Basement Complex zone (Gathright, personal communication). This arrangement is also more compatable with the groundwater discussion (Plate 2).

Elevations in the county range from a high in the western portion, two miles northeast of the Clarke-Fauquier-Loudoun county intersection, of 1920 feet above sea level, to 180 feet above sea

level in the eastern portion. The elevation at the primary watergap in the Blue Ridge belt at Snickers Gap (Route 7) is 1080 feet. The area between the two elevation extremes is characterized, going west to east, by high ridges and gradually descending, rolling hills.

The entire eastern half of Loudoun County lies in the Triassic Lowlands, a subdivision of the Piedmont. This area is characterized by flat-to-gently-rolling topography with several prominent northeast-southeast-trending ridges. The maximum relief of the ridges is approximately 100 feet. The cause of the oriented ridges is two-fold, being tectonic and erosional in origin.

# **Hydrology**

The Potomac River Basin contains all of Loudoun County, although the county is defined as part of sub-basin ll by the Virginia Department of Conservation and Economic Development, Division of Water Resources. The county is subdivided further into seven watershed basins (Plate 3). It should be noted that the preponderance of water falling on these watersheds flows either directly to the Potomac River (north and northeast borders), or into Fairfax County to become part of the Occoquan system. Streams which are found in Loudoun County are the North and South Forks of Catoctin Creek, North and South Forks of Beaverdam Creek, Little River, Goose Creek, Piney Run and Broad Run.

Three streamflow gauging stations are maintained in Loudoun County. Two of these are located on Goose Creek, whereas the third is on Catoctin Creek. The data from these are published by the U. S. Geological Survey in "Water Resources Data for Virginia." Flow data

		,

	{			
		•	•	,

from these stations for the fiscal years 1975, 1976, and 1977 are presented in Table 1.

Such data are of relevance, as the average discharge in cubic feet per second is related significantly to the drainage areas. The runoff of water from the watershed, translated into inches of water, can be compared to the rainfall for a similar period of time. Although the computations are complex, the runoff versus rainfall versus base flow also can be a gross indicator of the quantity of water available for recharge of the groundwater table. This recharge is, of course, not instantaneous but delayed due to the resistance of the media through which it moves. Other factors such as stream gradients, topography and soil types have strong effects on the recharge. See Plate 4 prepared by the Loudoun County office of the U. S. Department of Agriculture Soil Conservation Service.

The North Fork and South Fork of Catoctin Creek, Tuscarora Creek, Dry Mill Branch, Sycolin Creek, South Fork Sycolin Creek, Broad Run, Russell Branch, Beaverdam Run, Cabin Branch #1 and #2, South Fork Broad Run, and Sugarland Run were studied in detail for flood hazards by the U. S. Geological Survey, the Virginia Department of Highways and Transportation, the National Geodetic Survey, the U. S. Army Corps of Engineers, and the Division of State Planning and Community Affairs. The final report was presented in 1976 as a U. S. Department of Housing and Urban Development publication, Flood Insurance Study. The streams studied were perennial with distinct and variable channels. Broad Run and Sugarland Run are broad and flat, whereas the other channels are narrower and steeper. Although there

TABLE 1 LOUDOUN COUNTY AVERAGE DISCHARGE AT MAIN RIVER GAGING STATIONS PER FISCAL YEAR

June	69.0 73.8 193		64.5 59.6 161		30.4 16.4 48.2		State Highway
May	107 132 372		36.8 62.6 146		31.4 44.9 115.0		on State
Apr	102 131 395		129 160 468		193.0 170.0 485.0		bridge o
Mar	201 327 844		78.6 118 315		107.0 159.0 416.0	months	side of
Feb	120 176 484	June 76	94.6 157 393	June 77	**86.2 127.0 151.0	Averages based on incomplete months	« at Taylorstown, on right bank at upstream side of bridge  01638480)
Jan	112 140 402	y 75 -	268 346 985	y 76 -	* 111.0 98.5	ed on in	oank at
Dec 1975	138 269 908	1976 July 75 -	81.8 117 308	1977 July 76 -	73.9 128.0 286.0	ges base	right b
Nov	15.0 21.9 62.1		61.7 106 267	<b>,</b> —	74.7 122.0 288.0	۲ Avera	on, on
Oct	11.4 15.2 48.3		108 207 510		414.0 **272.1 *1715.0	φ *	Taylorsto 8480)
Sept	30.2 51.3 141		279 326 1005		28.8 28.0 84.7	rd by ic	reek at No. 0163
Aug	15.5 30.0 74.9		28.7 58.5 121		16.2 17.8 54.1	ed upwaı	Catoctin Creek 663 (USGS No.
July	15.6 37.4 85.5		137 171 339		38.6 37.4 130	Readings biased upward by ice	n 1 - Cat 663
Station	- 28		- 2 6		- Z E	* Read	Station 1 -

Station 2 - Goose Creek near Middleburg, on right bank 250 feet upstream from bridge on State Highway 611 (USGS No. 01643700)

Station 3 - Goose Creek near Leesburg, on left bank 400 feet upstream from bridge on State Highway 621 (USGS No. 01644000)

is little commercial, industrial or residential development in the floodplains, many roads and rail lines cross them. The cause of flooding in Loudoun County is almost entirely limited to local rainstorms and hurricanes. Large floods occurred in 1937, 1942, and 1972 (Tropical Storm Agnes). Thus far only one flood prevention lake has been formed; that being Chantilly Lake in the Broad Run watershed near Dulles Airport.

The aforementioned report presents data for predicted 10, 50, 100, and 500 year floods on all the mentioned streams. A presentation of the results of the flood study is beyond the bounds of this report. Detailed flood hazard maps are available of the Sycolin Creek, Tuscarora Creek, Broad Run - Sugarland Run areas, as noted in the bibliography.

Generally speaking, the quality of surface waters in Loudoun County is good, particularly the farther one is from the Potomac River. Catoctin Creek and Goose Creek are noted for good quality water. Even the Potomac, itself, in western Loudoun County, is fairly good. The quality of the lower Tuscarora, for the last 4.8 miles is poor. This latter condition is fairly true of most of the other streams; their quality degrades the nearer one goes toward the mouth. Water-borne sediment is also a problem in two areas of the County, with the Potomac #4 drainage area and Goose Creek basin losing up to 160 and 290 tons of soil material per square mile per year in 1968 (SWCB, Information Bull. 526, 1976).

### Climate

The climate of Loudoun is judged to be temperate, with an average

annual temperature range from  $-1^{\circ}\text{C}$  to  $23^{\circ}\text{C}$  ( $30^{\circ}\text{F}$  to  $73^{\circ}\text{F}$ ). Normal temperature extremes are from  $-24^{\circ}\text{C}$  to  $42^{\circ}\text{C}$  ( $-11^{\circ}\text{F}$  to  $107^{\circ}\text{F}$ ). The average annual rainfall is approximately 1.07 meters (42 inches) per year, although the year 1977 was exceptionally dry. The prevailing wind is from the southwest with seasonal winds from the northwest. Frequently the clash of the two air masses causing these winds produce heavy rains and subsequent flooding.

The tabulation presented below is taken from the Climatologic

Data assembled by the National Oceanic and Atmospheric Administration

for a period in excess of twenty years.

## AVERAGE MONTHLY TEMPERATURE OC (OF)

Jan. Sept. Feb. Mar. Apr. May June July Aug. Oct. Nov. Dec. 0 2 7 13 18 22 24 24 20 14 8 2 (55.0) (64.1) (72.4) (76.0) (75.3) (68.8) (57.7) (46.9) (36.0) (32.8) (36.1) (44.6) AVERAGE MONTHLY PRECIPITATION MILLIMETERS (INCHES)

Jan. Feb. Aug. Sept. Oct. Nov. Dec. Mar. Apr. May June July 59.9 60.4 95.3 101.6 119.4 86.1 93.2 72.6 71.6 61.5 85.6 81.0 (2.36)(2.38)(3.75)(3.19) (4.00) (4.70) (3.39) (3.67) (2.36) (2.32) (2.42) (3.37)

The recording station is located at Lincoln in west central Loudoun County. West of the station the winters are more severe and summers cooler than at the station. It should be noted that usually the hottest month is July, whereas the lowest temperatures are recorded in January. Further, the wettest month is normally June, with the driest months being the period September through November, and January through February.

Most groundwater is recharged during the fall and winter months

7 N.

2 UPLAND SOILS WITH SEASONAL HIGH WATER TABLES.

SOILS WITH HIGH SHRINK SWELL TYPE CLAY SUBSOILS.

WELL TO MODERATELY WELL, DEEP TO MODERATELY DEEP UPLAND SOILS WITH COARSE GRAINED SAPOLITES FROM CRYSTALLINE ROCKS.

UPLAND SOILS THAT ARE MODERATELY WELL TO WELL DRAINED, MODERATELY DEEP TO DEEP, DERIVED FROM FINE GRAINED CRYSTALLINE ROCKS.

SOILS DEVELOPED FROM OLD ALLUVIAL, COLLUVIAL, AND FLUVIAL DEPOSITS.

FLOOD PLAIN.

MORRISONVILI

WHEATLAND

BRITIAN

SOILS DEVELOPED FROM LIMESTONE CONGLOMERATE.

MODERATELY SHALLOW, WELL TO MODERATELY WELL DRAINED SOILS FROM TRIASSIC SHALES AND FINE GRAINED SANDSTONE

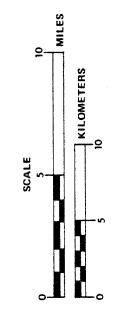
10 DEEP WELL DRAINED AND SHALLOW STEEP SOILS FROM SANDSTONE AND SANDSTONE CONGLOMERATE.

SYCOLIN

**EESBURG** 

LINCOLN

SNICKER GAP WOODBURN



BY: R. S. WEBER 1975 BASED ON "SOIL ASSOCIATION MAP OF LOUDOUN COUNTY" BY H. C. PORTER, JAN. 1952 PLATE NO. 4

•	•	•	

since vegetation utilizes the Spring and Summer rains. Previous severe droughts occurred in 1930-37, 1943-45, 1953-55, and 1964-65. Soils

As noted in the Loudoun County Soil Survey, 1969, the soils of the county are divided into 55 series and also eight miscellaneous land types. Such a diversity is explained by quoting <a href="#">The Soils</a>
<a href="#">That Support Us</a>. "First, there are those differences due to local variations in parent rock, slope and age. Secondly, there are those differences due to climate and vegetation. Then, there are some soils so young that no new features have developed, such as the fresh material just deposited by a stream, or the almost bare slopes of the mountains, or the dry, nearly sterile sand along the beach." (Kellogg, 1941).

Generally speaking, a given soil is divided into three zones:

a surface horizon, a subsoil and a parent material zone. Another

way of considering the soil is to think of three horizons overlying

the rock, or parent material. Within these three basic horizons,

and often overlying them, are other zones. Such a situation expressed

graphically, for a mild temperate climate with moderate rainfall,

would be as shown in Figure 1.

The A horizon is characterized by the chemical alternation of the soil material in acid, reducing environment. Clays of the kaolinite family, soluble carbonates and simisoluble iron minerals, if produced, are leached downward by moisture, leaving the horizon silica-rich. Therefore, the center portion of the zone is usually sandy and light colored. The upper portion is dark-colored from organic material.

Horizon 0		Description	
		Organic litter, humus	
	A <sub>1</sub>	Organic Colloids & Minerals	hing n
A	A <sub>2</sub>	Light Colored leached	Zone of leaching or eluviation
	A <sub>3</sub>	Leached, but transitional	
	B <sub>1</sub>	Accumulation transitional	ation
В	B <sub>2</sub>	Accumulation, clay formation, deep color	Zone of accumulation or illuviation
	В3	Transition, more color than C, carbonate accumulation	Zone of or illu
С		Silcia, Carbonate, Sulfate accumulation, slight weathering	
D or R		No Alternation	

Figure 1 Idealized Soil Profile in a region of mild temperature and moderate rainfall. (After Sowers & Sowers, 1970)

The B horizon is thicker than A and contains an accumulation of clay materials, iron, and carbonates. The  $B_1$  layer can be partially cemented and deeply colored, whereas the  $B_3$  layer is lighter in color.

Below the two true soil horizons is the C layer of slightly weathered, or crumbly bedrock. If an ancient, unrelated soil, a paleosoil, is present, it is called the R horizon; otherwise the base layer, if rock, is the parent material, D.

According to recent thinking there are in Loudoun County ten major groupings based on major, similar, soil properties. These have features affecting both urban and rural use. It is stressed that the coding indicating certain properties is not all inclusive and does not apply to each individual soil mapped within an association. That is the type of detailed information which should be obtained from the detailed soils maps for a given tract.

The ten soil groupings are listed in Table 2 and briefly described.

It is not the purpose of this publication to describe and define all the soils of Loudoun County. Such data is available in the Soil Survey series 1951, no. 8, issued September 1960 by the U. S. Department of Agriculture. This publication and other more recent ones are available through the local Soil Conservation Service office.

The primary purpose for mentioning the soils is that survey publications contain detailed cross-sections of all the soils within the county. These cross-sections provide data on the thickness of the soil layers and the internal-drainage characteristics. These are an important factor in water-well productivity, particularly in areas underlain by crystalline bedrock. The internal drainage characteristics affect the rate at which water infiltrates the soil and gives an indication of the recharge rates. Soil texture is the major determining factor affecting internal drainage. A coarse-

## TABLE 2

## LOUDOUN COUNTY

## SOIL GROUPINGS

Physical Characteristics Found on steep slopes Many rock outcrops Thin soils

Most found in Triassic Lowland Water table highest from January to Apřil

Upland soils with seasonal

\_

high water tables

from crystalline rocks

Steep & shallow soils

Soil Group & Name

Usually overlie diabase and basaltic rocks of the Triassic Lowland

Overlie mostly basaltic-type Located in western-most Piedmont Uplands and Middleburg rocks

coarse-grained saprolites

Deep upland soils with

 $\geq$ 

from crystalline rocks

Largest single unit overlying granitic-types with minor greenstone. Highly variable in depth, clay content and crystalline rocks, mostly grainsize

grained crystalline rocks

Upland soils from fine-

Water-bearing Characteristics for Overburden Wells Locally productive immediately over Not dependable bedrock

Locally productive for shallow wells Easily contaminated

Well yields would be seasonally controlled Jsually reliable for single-family Good recharge potential dwellings

Variable characteristics prevent any overall predictions

Soils with high shrink-

ΠΙ

swell clay subsoils

(cont.)	COUNTY	UPINGS
2	NDOOL	GROUP
TABLE	COUDC	SOIL

LOOD	SOIL	

for	
Characteristics	
teris	
ıarac	
	We]]s
eari	den,
ter-bearing	rerburden
Wat	0

Alluvial, Colluvial, and Fluvial Soils

Soil Group & Name

Found at the base of mountains particularly Lovettsville and

Lucketts areas

Physical Characteristics

depths or remains as a local reservoir Water usually passes quickly to lower perched on a clay bed Usually coarse, cobbly deposits, but can contain clay lens

> Flood Plain Soils ٨II

Highly variable composition though usually contains silt and clay size material mixed with rock beds

Generally productive, however restricted by Flood Plain Insurance Regulation

Soils from limestone Conglomerate VIII

Surface wells not usually feasible Very complex and frequently

> Moderately shallow fairly well drained soils of Triassic Clastics  $\simeq$

Triassic Basin overlies

overlie sinkholes

shale and fine grained

sandstone

Usually less than 660 mm (26 ft) thick

Not practical for shallow wells Mostly wide-spread soil of

> shallow steep soils of Triassic conglomerates Deep well drained and

Limited to areas northeast and southeast of Oatlands Characterized as cobbly and possessing rapid internal drainage

Well production would be most dependable at soil-rock interface textured soil will retain and recharge water more rapidly than a fine-textured clay, or clay loam. An additional benefit of the official soil survey publication is the engineering chart, which summarizes the texture, permeability, depth to bedrock, depth to water table and other characteristics.

## Vegetation

At present only about 16 percent of the land remains in forests. Little of this is virgin growth. The predominant timber is hardwood, consisting of oak, hickory, maple, beech, yellow poplar, black locust, walnut, sassafras, dogwood, and persimmon. Weeds and deliberate pasture plants grow generally well. Additionally, the bushes and shrubs of the forest areas are fairly prolific. Needless to say, vegetated areas act as better watershed and recharge zones than do highways and urban zones.

### CHAPTER III

### HYDROGEOLOGY

## Geological Setting

Occupying, as it does, two physiographic provinces composed of four sub-provinces, the geology of Loudoun County is very diverse and complex. Therefore, the hydrologic system developed on and by the geology is not simple. The original geologic units frequently have been altered by weathering, tectonics, vegetation and land use. The rocks and rock materials, themselves, are assigned to three basic groups based upon origin. The three origins are sedimentation, igneous emplacement, and metamorphic processes. This gives rise to the groups - sedimentary, igneous, and metamorphic rocks (Plate 5).

The oldest rocks found in the county are the granite and granite-like rocks of the Marshall Formation and granulite gneiss to the Southwest. Though granite-like all these rocks are gneissic in nature. Above this group are other metamorphic rocks of sedimentary origin. These rocks include sandstones, boulder beds, conglomerates and recrystallized limestone. Above this complex were deposited many layers of basaltic lava flows. These also have been metamorphosed, with three easily recognized facies being present.

The metabasalt, called the Catoctin Formation, covers a wide area of the county and is the primary ridge-former of the Blue Ridge Mountains. Overlying the Catoctin is a narrow belt of up to 500 feet-thick quartzites, interbedded phyllites and cross-bedded conglomerates.

Throughout most of the county, the Bull Run fault marks an

abrupt change from Cambrian (?) age, 570 M.Y.B.P., to Triassic age 190-225 M.Y.B.P., sedimentary rocks. However, in the extreme north central and south central portions of the county wedges of metamorphic rock border the lowest edge of the fault. Both of these wedges, approximately 122-152 meters (440-500 feet) wide at the widest points, taper to zero several miles short of the Washington and Old Dominion Railroad tracks west of Leesburg.

The northern wedge contains metamorphosed sediments of the Hampton (Harpers) Formation and the Erwin (Antietam) Formation of quartzite and silty phyllite. The southern wedge, on the other hand, is almost entirely quartzite of the Weverton Formation, with a very thin band of Loudoun Formation phyllite on the west (Gathright, personal communication).

As stated previously, the Bull Run fault forms the physiographic boundary and marks the western border of the Triassic lowlands portion of the Piedmont Province (Plate 2). It is thought that the fault has a vertical displacement of as much as 3048 meters (10,000 feet) with the eastern block being downthrown (Parker, 1968). This block is comprised of sedimentary rocks which have locally been intruded by igneous rocks. The sedimentary rocks are grouped into four basic types: conglomerate, both limestone and quartz; sandstone; shale or siltstone; and pyroclastic rock.

Several theories have been put forth to explain the origin and subsequent geologic history of the Blue Ridge - Piedmont section of the United States. But regardless of theory, there are only a few places where stratigraphic units can be correlated from one side of the Blue Ridge to the other. Thus, establishing a correlation between

* *		

the well-known miogeosynclinal rocks west of the Blue Ridge with the very-poorly-known eugeosynclinal (?) rocks of the Piedmont has become a major problem in Appalachian geology studies.

Some authors feel that the Blue Ridge occupies the site of a welt, or "median geanticline", that persisted throughout nearly 300 million years (Reed, 1970). Related to this idea, Drake, Ewing, and Sutton proposed, in 1959, that the Blue Ridge, deeply buried at the time, was the dividing zone for two separate basins of sedimentary deposition; the western basin being a shallow shelf trough, while the eastern trough was deep receiving potentially over 6100 meters (20,000 feet) of material. Basement uplift later exposed the entire eastern continental margin to erosion and produced today's Blue Ridge.

Others authors, notably Dewey and Bird (1970a, 1970b) say that the Blue Ridge, being basaltic rock, represents the site of upwelled, oceanic crust material caused by continental collision. The collision in this case was between the North American and African plates. Thus the Piedmont Province is really a residual portion of the African continental glued onto the North American continent.

In any event, the oldest rocks, the granitic gneisses, are thought to be 1000 to 1100 million years old. These rocks were uplifted, exposed to erosion, and locally overlain with variable thickness of boulders and sediments.

Beginning in late Precambrian time lavas and tuffs of the Catoctin system were deposited over the basement rock and local sediments. Further sediments were deposited between the lava flows.

Feeder dikes and sills of gabbro penetrated underlying rocks at this time (Brown, 1970).

The weight of the load of sediments caused subsidence in the eastern trough and to a lesser extent the western trough. Some local uplift took place of sufficient duration for a soil to form before reburial; thus producing in local areas the Loudoun Formation. Above the Catoctin Formation, and local Loudoun, next was deposited the quartz sand of the present Weverton Formation, followed by mud, silt, and sand forming the upper Weverton and Hampton (Harpers) Formations. The last-recorded period of deposition was the sandy material of the Erwin (Antietam) Formation. If any later deposition took place, it has been removed by later erosion.

In later Paleozoic time, horizontal forces from the southeast (continental collision?) produced compressional strain in the rock units. More competent rocks were fractured and faulted, while less competent rocks flowed. The forces also caused metamorphism by heat and pressure, which changed many of the sedimentary units: i.e., sand, limestone, silt, etc., into quartzite, marble, and slate respectively.

During Triassic time, normal tension faults developed, with down-dropping on the east side. Associated with this cracking due to tensional stress, more igneous material, mostly basaltic in nature, was emplaced. The ancestral Triassic Basin, to the east of Bull Run fault, began receiving clastic sediments. The granites were formed by deep intrusion of molten rock into already-established beds. These subsequently have been exposed by heavy erosion. Such erosional forces are still at work producing the landforms of today.

## Geologic Formations

A simplified geologic map of Loudoun County is shown in Plate 5.

Table 3 lists the rock formations. It should be noted that the column begins with the oldest at the bottom and becomes younger upward; in the order that the rocks would be encountered by a driller. The water-bearing properties have also been noted in general terms. The exact age relationships are also indefinite in some cases, as much more geologic work is needed. A discussion of the groundwater characteristics of the formations according to lithology and physiographic province is presented further in the text.

The Marshall Formation is the oldest recognizable unit in the study area. The unit has been called the Basement Complex (Bloomer and Werner, 1955), Virginia Blue Ridge Complex (Brown, 1958) and various formational names. These rocks are found in a wide belt 5.6-10.5 kilometers (3.5-6.5 miles) wide in the western half of the county, bounded on both sides by the Catoctin Greenstone. Several types of intermediate and granitic gneisses are widespread in the belt. These generally are considered by the geologic profession to have been of igneous origin, later metamorphosed. Some authors have reported what they believe to be rocks of sedimentary origin intruded by granites (Jonas and Stose, 1939). Parker (1968) considered the eastern side of the belt as being the shallower edge of the Marshall and reported it as being characterized by smaller crystal sizes. In the western, deeper side, he reported feldspar phenocrysts up to 50 millimeters (2 inches) across. The coarser texture and different associated minerals, i.e. garnets and biotite, suggests a higher grade of metamorphism for the western edge. An extensive quartz vein

TABLE 3

# GEOLOGIC FORMATIONS AND THEIR WATER-BEARING PROPERTIES

Water-Bearing Properties	very poor aquifer, dry holes common.	very good aquifers; deep drilling (in excess of 500 ft.) feasible; best wells usually less deep.	Good water bearing units at depths less than 400 feet.
Lithology	Dark, blue-grey- to-black igneous rock midway between basalt and gabbro in grain texture.	Conglomerate, sand- stone, shale, basalt pyroclastics and zones of contact metamorphism. Most sedimentary units are red in color. Large local deposits of lime- stone conglomerate around Leesburg.	Blue slaty, crystal- line limestone
Physiographic Occurrence	Triassic Lowlands	Triassic	Layered Clastics
Rock Unit and Symbol 1	Diabase (basalt) (TRB)	Newark Group (TRLC) (TRBF) (TRBB) (TRMS) (TRMR)	Frederick Limestone (EF?)
Age (Geologic Period)	Triassic	Triassic	Ordovician - Cambrian

TABLE 3 cont.

## GEOLOGIC FORMATIONS AND THEIR WATER-BEARING PROPERTIES

Water-Bearing Properties	Generally not water- bearing unless faulted.	Generally a low water producer requiring inter- mediate depth wells - 200-400 ft.	Poor aquifer
Lithology	Predominately an interbedded sequence of white, fine-grained quartzite with minor darker phyllites or sericitic quartzite.	Light grey to greenish-grey phyllite inter-bedded with layers of fine-grained quartzite, locally calcareous, some graywacke and medium to thick-bedded, coarsegrained quartz sandstone.	Light grey-to-white, fine-grained, finely-laminated-to-bonded quartzite. Some quartz-pebble conglomerate, pebbly quartz sandstone, and very sericitic sandstone, phyllite and sandy
Physiographic Occurrence	Formation Layered Clastics	Layered Clastics	Layered Clastics
Rock Unit and Symbol 1	Erwin (Antietam) (EER)	Hampton (Harpers) Formation (£H)	Weverton Formation (EW)
Age (Geologic Period)	Cambrian	Cambrian	Cambrian

TABLE 3 cont.

# GEOLOGIC FORMATIONS AND THEIR WATER-BEARING PROPERTIES

Water-Bearing Properties	Only locally occurring, generally poor.	Generally considered poor water-bearing. Yield and depth in the low (shallow) to moderate range.	Generally poor aquifer due to shallowness.
Lithology	Light-purple, tuff- aceous phyllites, unit is not mappable, discontinued as a stratigraphic term. (Gathright and Nystrom, 1974, p. 15) (not shown on map)	Interbedded, reddish- purple, tuffaceous phyllite; dark-green metabasalt; white, pink and tan rhyolitic metatuff, epidotized quartzose meta-sediments; metabasalt, basaltic flow breccias, epidosite and purple amygdaloidal slate.	Discontinous, beds of shallow, massive-to-medium-bedded, white, blue, gray, green, dull-to-poorly-lustered marble. (not shown on map)
Physiographic Occurrence	Layered Clastics	Layered Volcanics	Layered Clastics
Rock Unit and Symbol	Loudoun Formation (?) (£L)	Catoctin Formation (CPCC)	Everona (?) Formation
Age (Geologic Period)	Cambrian	Upper Pre- Cambrian	Upper Pre- Cambrian

TABLE 3 cont.

## GEOLOGIC FORMATIONS AND THEIR WATER-BEARING PROPERTIES

Age       Rock Unit       Physiographic       Lithology         Upper Pre-Cambrian       Swift Run Layered Clastics and phyllites. (CPCS)       Heterogenous unit of metagraywacke, metagray	Water-Bearing Properties	Generally good for slow demand use.	Dependable aquifer at depths less than 200 but only moderate quantity of water.
Rock Unit and Symbol Swift Run Formation (CPCS)  Marshall Formation (PCM) (PCG)	Lithology	Heterogenous unit of metagraywacke, metagraywacke, metagrakose, and phyllites. Overall texture ranges from gneissic boulder beds upward through cobble beds and arkoses to cross-bedded, mediumgrained arkose and platybedded sandstone.	Light-to-dark gray, medium-grained granitic rock, including quartz monzonite syenitic granite and augen gneiss. Felsic and mafic dike intrusions.
Derriod an	Physiographic Occurrence	Layered	Basement Complex
Age (Geologic Period Upper Pre- Cambrian Pre-Cambrian	Rock Unit and Symbol	Swift Run Formation (CPCS)	Marshall Formation (PCM) (PCG)
	Age (Geologic Period	Upper Pre- Cambrian	Pre-Cambrian

 $<sup>^{</sup>m 1}$  Symbols correspond to aquifer listing in Appendix B

Sources: Cady (1938); Toewe (1966); DeKay in Lukert and Nuckols (1974); Gathright and Nystrom (1974); Mack (1965); Parker (1968); and Virginia State Water Control Board - NRO

system occurs on the eastern half and leaves masses of blue quartz when weathered.

In the case of a eroded anticline, such as the Blue Ridge area, the rock units decrease in age from the core outward. Thus, the next unit encountered on each side of the Marshall Formation is the Swift Run and Catoctin Formations, depending on local conditions. Late Precambrian erosion of the gneissic and grantic basement rocks produced a relief of several hundred feet. Into these drainage features were deposited boulders and cobbles of the basement rocks, intermingled with finer mineral products of mass-wasting. The depositional pattern was discontinuous, yielding here some deep beds of one type rock and in other areas shallower beds. The rock types grade upward from boulder beds to cobble-conglomerate to coarsegrained arkose and platy-bedded, fine-grained feldspar-bearing sandstone. The beds are often lenticular and of limited lateral extent. Sericite and quartz form 65-90 percent of the groundmass. Other minerals such as chlorite, epidote, and iron compounds occur locally.

At the top of the phyllitic arkoses, where they grade into phyllites, there are zones of bluish-white, recrystallized limestone (marble) lenses, typically shallow in depth and up to 30 meters (100 feet) long (Parker, 1968). These were called the Everona Formation by Tinsely Mack in 1965. They appear to be enclosed in the upper beds of the Swift Run Formation. Their thickness ranges from 4 to 29 meters (14 to 95 feet), whereas the colors and textures range from white, gray, light blue, and light green, with fine to medium crystals.

Some talc, graphite and mica occur as impurities in individual outcrops. Recent work has determined that the limestone of this type

found in Loudoun County are not typical of the Everona, therefore suggesting the use of the (?) after the term (Gathright, personal communication).

The bulk of the western half of Loudoun County is composed of two bands, the eastern being the wider, of the Catoctin Formation. This unit is composed almost completely of greenschist-facies, metamorphosed, basaltic rock. The most typical form is a dark-gray-to-blackish-green, massive, lava flow with fractures. A pistachio-green, epidosite form is the next most common form. The third variety is a chloritic phyllite which apparently marks zones of maximum stress-relief during diastrophism (Parker, 1968). Between the flows are found beds of phyllites, rhyolitic metatuff and metasediments, and purple slates. The latter in an amygdaloidal form was previously called the Loudoun Formation.

The next three rock units are lumped together as the Chilhowee Group and are detrital marine sediments which have undergone low-grade metamorphism. The lowest unit, approximately 137-183 meters thick (450-600 feet), is the Weverton Formation. The formation occupies a position along the West Virginia border of Loudoun County and another linear slice adjacent to the Bull Run Fault. The most common constituent of the Weverton is a quartzite, with lesser quantities of phyllite, quartz pebble conglomerate, sandstone, and sandy phyllite. Generally, the quartzite is impermeable.

Stratigraphically above the Weverton Formation is the Hampton (Harpers) Formation of laminated-to-thick-bedded, dark-grayish-green, very-fine-grained, phyllitic, feldspathic graywacke with ledge-forming, medium-to-thick-bedded, coarse-grained quartz sandstone and quartzite.

The unit is reported to be up to 61 meters (200 feet) thick west of the Bull Run Fault.

Continuing upward from the Hampton Formation is found the Erwin (Antietam) Formation of white, fine-grained quartzite, silty phyllite, and sericitic quartzite. From the Potomac River the unit extends southwestward for about 7.2 kilometers (4.5 miles) before tapering downward into the Bull Run Fault.

About 2.4 to 3.2 kilometers (1 1/2 to 2 miles) north of Lucketts and west of the fault a small mass of lower Ordovician Frederick limestone is found (Cady, 1938). This thin exposure consists of a blue, slaty, crystalline limestone. The depth of the formation is thought to be between 30-61 meters (100-200 feet). Although called Ordovician in age by Cady, the Virginia Division of Mineral Resources has labeled it as Cambrian. This is the reason for the question mark in the Explanation of the geologic map.

The Newark Group comprises the suite of sedimentary rocks,

Triassic in age, which lies east of the Bull Run Fault and comprises
almost the entire half of Loudoun County. The Group has been divided
by K. Y. Lee (United States Geological Survey, 1977) into the Manassas
Sandstone, Balls Bluff Siltstone and Bull Run Formation. The Manassas
has a Reston member and a sandstone member. The Bull Run is also
further divided into the Leesburg Limestone Conglomerate member and
Basaltic-flow bearing clastics member (Plate 5). These rocks,
predominately red to maroon in color, include conglomerates, sandstone, siltstone, and shale. All the various units are interbedded
vertically and interfinger laterally. Additionally, several large
bodies of igneous rock, primarily diabase (bluestone), have been

injected into the sedimentary units. Other igneous rocks noted are syenite and basalt. The syenite resembles granite, but lacks the normal quartz content. Most of the igneous rocks are found as dikes, sills or stocks, although many lava flows and volcanic ash deposits are also present.

The limestone conglomerate outcrops west of Leesburg in a large north-south trending deposit. It interfingers with sandstone, shale and pyroclastics south and east of Leesburg. The limestone pebbles are from 6.35 millimeters (0.25 inches) up to several feet in diameter, all in a red matrix of quartz, feldspar, calcite, mica, chlorite, and clay (Toewe, 1966). The majority of pebbles are calcite or dolomite, whereas about 25 percent are quartz, and the remainder metabasalt, schist, feldspar, phyllite, and shale pebbles.

The quartz conglomerate is exposed throughout a large portion of the Southern part of the county, particularly around Leesburg. Here the pebbles are 45-55 percent quartzitic material, 6.35 millimeters (0.25 inches) to 152 millimeters (6.0 inches) in diameter. Such deposits are up to 183 meters (600 feet) in thickness.

The shales and sandstones owe their color to varied proportions of iron oxide, both as a constituent and as grain coatings. The shales range from soft to extremely brittle. The sandstones frequently contain high enough proportions of feldspar to be counted as an arkose. Colors noted are pink, gray, red, maroon, brown, and green. Grain textures range from medium to coarse. The Manassas and Bull Run units represent fan-shaped river or stream deposits, whereas the Balls Bluff was deposited chiefly in lakes. It should be mentioned that the latter beds of siltstone are frequently calcareous.

In the area between Arcola and Conklin and along Virginia routes 659 and 772 there are several zones of "baked" shale. This dark, dense rock, called hornfels, forms adjacent to igneous intrusions where country rock is cooked by the high temperatures. The metamorphosed rock may resemble a basalt. The zone of affected rock can extend up to 0.8 kilometer (0.5 mile) from the intrusion and exhibit transitional metamorphism from heat source outward.

## Geologic Structure

In the most simplified terms the structural geology of Loudoun County may be thought of as a large, eroded, upward-arched fold of rocks (western part), bounded by a deep normal fault and flanked to the east by a large basin of sedimentary, igneous and volcanic debrisorigin rocks. The regional trend of all the structures is close to  $N20^{\circ}E$ .

The Blue Ridge portion of the county may be thought of as one large upwarp of the earth's material 12.9 kilometers (8 miles) wide upon which smaller folds 20 meters to 200 meters (66 to 650 feet) were superimposed. There were several periods in geologic history when strong horizontal, compressive forces further upwarped the surface and exposed the whole area to erosion.

The Triassic Basin, part of the larger Culpeper Basin, was uplifted at its northern end and downwarped in the south. It was later tilted westward and north and downdropped along the Bull Run fault. The diabase intrusions occurred during the dropping (Lee, 1976). The Bull Run Fault, itself, extends north and south out of Loudoun County into Maryland and southern Orange County, respectively, a distance in excess of 145 kilometers (90 miles). The fault-plane dips

southeast at on an angle of  $30^{\circ}-50^{\circ}$  (Keith, 1893). Other normal faults are also present in the county, but most are roughly perpendicular to the Bull Run. Most of these occurred after the diabase intrusions.

The tectonic forces to which the area was exposed also caused the regional joint system. Three joint sets are evident in the Triassic sedimentary deposits. Predominately these are east-west, nearly vertical to vertical, and N50 $^{\circ}$ E, with dips 70 $^{\circ}$ SE to vertical. In the rocks on the west side of the fault the dominant joint sets are N10 $^{\circ}$ S, N80 $^{\circ}$ E, and N45 $^{\circ}$ W, with all dips being nearly vertical. Such joints can be the sites for small-volume water supplies.

The basin sediments are thicker on the west side and thin-to-a-feather-edge on the east. The angle of inclination is  $15^{\circ}-20^{\circ}$  to the northwest, with the greatest dips in the west.

## Groundwater

Groundwater is an integral part of the earth's water circulation system, which is called the hydrologic cycle (Plate 6). Groundwater originates from precipitation in the form of rainfall or melting snow. That portion of the precipitation available to groundwater recharge is equal to the total precipitation minus evaporation, transpiration by growing plants, runoff to streams, and the amount held as soil mixture. Some of the water that infiltrates through the soil zone is held by capillary forces above the water table; the remainder moves downward by the force of gravity and replenishes the groundwater, causing the water table to rise. Therefore, countrary to popular opinion, each rain does not recharge the groundwater table.

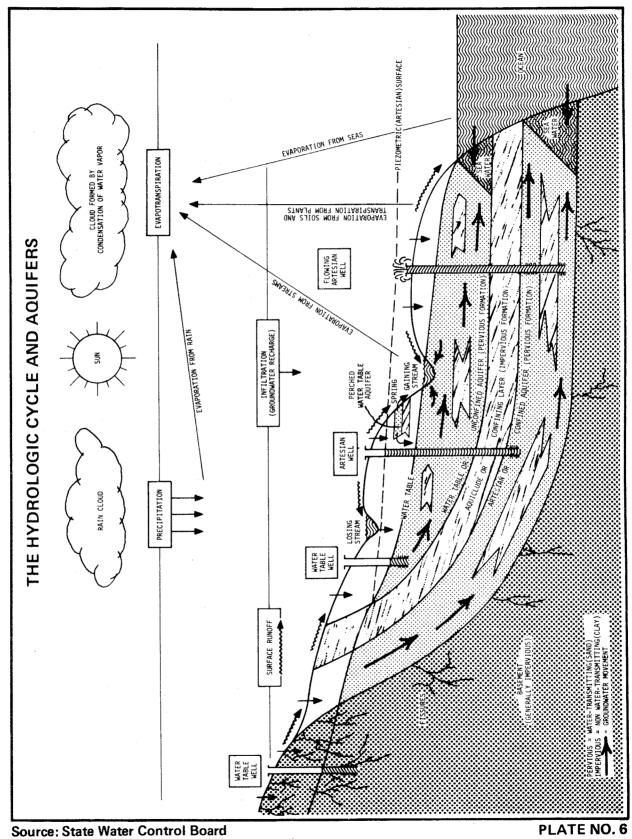
	Ground Level
Zone of aeration (indefinite depth)	Discontinuous moisture zone Moisture films not connected
Capillary Fringe Zone	Soil partially saturated, moisture continuous
(indefinite_depth)	
Capillary Saturation Zone	Soil saturated, moisture continuous
∇ Water Table	

Figure 2 Soil Moisture Distribution

The surface of the water table is rarely flat, but has undulations usually conforming to the topography. The water table is higher under hills than under valleys. However, the relief of the water table surface is more subdued than the topographic expression. Therefore, the depth to the water table is greater under a hill than under a valley.

The differences in elevation of the water table and the force of gravity cause movement of the groundwater. This movement often results in the natural discharge of groundwater through springs. Springs or seeps occur where the water table intersects the land surface, usually on the sides of slopes and in valleys. Thus, high topographic areas are commonly groundwater recharge zones, and low topographic areas are commonly groundwater discharge zones.

The dominant factor that influences the occurrence of subsurface water is the geology of the area. This is particularly true where



		•
		•
		i
		•

the underlying solid rocks have a very low permeability. In these cases subsurface fluids are transmitted mainly through fractures and solution openings. The former occur due to deformation forces and the latter due to decomposition of mineral constituents. Such voids are referred to as secondary permeability and are directly related to the structure and lithology of the rock units present (DeKay, 1976).

Groundwater does not flow in "underground streams" except in limestone or dolomite caverns, which are probably limited to the Leesburg area. Instead, groundwater normally flows through the interconnected pore spaces in sediment and fractures in rock. The rate of movement ranges from a few inches per year to a few feet per day.

Groundwater movement in crystalline rock is affected by structures in the bedrock, particularly fractures such as faults and joints. The direction of movement is affected by the alignment and inclination of the fractures. The rate of movement is influenced by the number and size of the fractures, and the depth of movement is determined by the depth to which the fractures extend.

at wells. The pumping of a well will cause the water level in the vicinity of the well to drop, resulting in movement of groundwater towards the well. The vertical distance that the water level in the well drops is termed the drawdown. The drained space between the original water level (static water level) and the lowered water level (pumping water level) outside the well is roughly cone shaped, and is termed the cone of depression. The cone of depression spreads outward from the well as it is pumped. The cone of depression eventually will stabilize at some distance for the well if the with-

drawal rate is not excessive. If two or more wells are placed too close to each other, the cones of depression of each will intersect and may result in excessive drawdowns, decreased yields, and possibly the "drying-up" of one well.

## <u>Aquifers</u>

An aquifer is a geologic formation that is capable of containing and transmitting groundwater. Aquifers may be comprised of unconsolidated sediments or of fractured, consolidated rocks. Groundwater is contained in the pore spaces of sediments, such as sand or gravel, and in fractures or cracks in crystalline rock, such as granite.

Aquifers may be of two types: (1) unconfined or water-table aquifers and (2) confined or artesian aquifers. Both types are present in Loudoun County and examples are illustrated in Plate 6.

In an unconfined aquifer, the water table is the upper surface of the zone of saturation. The surface of the water table is at atmospheric pressure, whereas water below this surface is at a higher pressure due to the weight of the overlying water. In an unconfined aquifer, water is able to infiltrate downward through the zone of aeration and reach the water table.

In an artesian aquifer, the groundwater is separated from the zone of aeartion by a confining layer, such as an impermeable clay bed. Groundwater in a confined aquifer is under pressure in excess of atmospheric pressure. When the confining layer is penetrated by a drill bit, the groundwater quickly rises in the bore hole. The water level in an artesian well stands above the aquifer that the well taps. If the artesian pressure is sufficiently high, the water may flow freely at the surface. However, a well does not have to

flow freely for an artesian condition to exist. Cady (1933, 1938) reported there were few to no artesian systems in Loudoun County and what few there were would be limited to the southern part of the basin near the fault line.

The depth to the static water level of an aquifer can be measured in any well that is not being pumped. If the well is in a confined aquifer and the upper water table is cased-off, the measured water level is termed the artesian surface (or potentio-metric surface). In a well tapping an unconfined aquifer in which water from the upper zone is allowed to enter the well, the measured water level is termed the water table. The depth to the groundwater in either case will depend on several factors including: the geologic formation in which the well is drilled, the confining condition of the aquifer, the topographic position of the well site, the time of year the measurement is taken (water levels fluctuate seasonally), and the proximity to pumping wells.

Groundwater sometimes flows for many miles in extended aquifers, but frequently, especially for unconfined aquifers, it has partly or totally originated from surface infiltration within a radius of a few miles of its occurrence. The recharge area for most water-table wells is the immediate area around the well for a radius of several hundred feet. The Triassic Basin tends to act as its own recharge area, with nearly all the absorbed water being used through the wells drilled there or going away as the base flow of the streams. Therefore, it is important what contamination sources are located within the basin. Recharge of an artesian aquifer may occur from lateral migration from the outcrop area of the aquifer or from vertical

leakage through semipervious confining beds.

## Basement Complex - Blue Ridge Province

The Basement Complex, consisting of the Marshall Formation and other granite-like rocks, is not generally thought of as a good water-bearing complex. Where massive, the rock is practically impermeable except for the faults and joints. Toward the eastern edge of the outcrop area, however, the schistose planes are more open due to tectonics and weathering. Here water for domestic use is generally abundant. The yields decrease below 120 meters (400 feet) as the cracks and joints become sealed by pressure. Cady (1933) reported that below 60 meters (200 feet) a 19 lpm (5 gpm) increase in yield could be expected for every 30 meters (100 feet) of additional depth. Beyond 120 meters he made no predictions.

The controlling factors of many well sites in western Loudoun County are the thickness of the soil overburden, the porosity and permeability of the overburden, the topographic position of the well site, degree of fracturing of bedrock, occurrence of mineralized zones, and the lithology of the parent rock. The first three items affect the storage capacity and rate of recharge of the rock. This is particularly true in the upper zone of the bedrock and in the presence of joints.

The topographic position of the well site is very important, as wells drilled in areas of low elevations statistically have better yields. This is probably due to such areas having buried faults, joint lines, soluable bedrock, and decomposed strata. More to the point, low areas are receiving recharge from higher elevations within the same rock units.

Drilling in the Marshall Formation will, except for local areas, be difficult, accompanied by a high rate of bit wear. Acceptable water frequently will be found just at, or slightly below, the zone of decomposed bedrock. The volume will usually suffice for domestic use. Special screening precautions should be taken when driving the casing in order to take advantage of this water.

Frequently concentrated zones of minerals will produce better water sources than the parent rock. This is true in cases where layers of brittle quartzite are enclosed in granites or basalts; or where calcitic beds are found in between basalt flows or dense conglomerates.

## Layered Clastics

This suite of rocks occupies the smallest area, yet contains a great diversity in formations and lithologies. The seven formations include boulder and cobble beds, arkose, phyllite, graywacke, quartz-pebble conglomerate, sandstone, quartzite, limestone, and marble. The possibilities of finding useable quantities of water in the various units are lithologically and tectonically controlled. The lithology effects the water holding and storage capacities chemically, whereas tectoncis caused mechanical breakage.

The Swift Run occurs in outcrop bands and several irregularshaped areas. Texturally the unit ranges from a boulder bed to medium-grained sandstone.

The banded areas should be considered as poor-quality wateryielding beds, although the irregular bodies in the south central portion of the county have had wells producing in the 25-100 gpm range. The unit probably can be over-pumped easily. Some of the best-quality water could be expected from the so-called Everona (?) Formation. However, due to the shallowness and erratic occurrence of the marblelike limestones, it cannot be relied upon. Those wishing to exploit the unit should refer to Tinsely Mack's report already mentioned.

As in the case of the Everona (?) Formation, the Loudoun should not be relied upon. Its extremely-limited outcrop and shallowness almost assure that any wells located in it will be of the overburden type or of low output.

The next three formations of the Layered Clastics, the Weverton, Harpers and Erwin, can be lumped together roughly for the discussion purposes. The only unit which has been identified definitely in well cuttings is the Weverton. Indications are, however, that wells have been made in all of the formations. Table 2 presents the water-bearing capabilities of all three. Further discussion of the individual units can be found in Table 5 of Chapter 5.

The last unit of the layered clastics, the Frederick Limestone (Cady, 1938), is limited to a small section north of and adjacent to Route 15, east of Point of Rocks. The thickness of this unit is unknown, as all the confirmed wells in the Frederick are less than 200 feet in depth. The yields have been reported as unknown, but adequate for a family dwelling, up to 10 gpm. It is speculated that this unit should be capable of yields in the 25 gpm-and-above range. Layered Volcanics

Generally speaking the layered basalt flows, with interbedded metasediments, cannot be considered to be a good source of groundwater supplies. It was noted that prior to the advent of modern drilling

technology nearly all the wells in the Catoctin formation were shallow, limited to the overburden. As in many other areas, the topographic position of the well site is of great importance. Low areas, particularly if linear in configuration, could be surface expressions of deep-seated fractures, and fractures serve as better water sources. Such fractures are more numerous nearer the surface and decrease with depth.

The contact surface between wet sediments and volcanic flows or quartz zones are usually good water sources in the Catoctin. However, the chances of predicting such occurrences are found in areas mapped as Marshall Formation. The lines between these units are potentially-better water zones than either formation. Yields from the Catoctin area have almost always been in the poor-to-infrequent-fair range [less than 38 to 76 lpm (10 to 20 gpm)].

# Triassic Lowlands

The sedimentary rocks of this large section of the County are the best water sources. Although wells have been drilled to depths of almost 300 meters (1000 feet), and a percentage of wells in the 182-243 meters (600-800 feet) range have been dry or inadequate, it is generally a reliable aquifer. This is due to the geologic controls: i.e., depth, dip, degree of fracturing, lithologies, etc. Nearly the entire basin can be summarized as containing potentially-fair-to-good resources. One area in particular around Leesburg is evaluated as very good, 378.5 to 946 lpm (100-250 gpm) wells being common.

One groundwater danger which exists in the Triassic, particularly around Leesburg, however, is the practice of overpumping. This act, committed in order to obtain, rapid, high volumes of water, can lead

to a loss of hydraulic lift for the overlying rocks and soils above the water body. Such a loss ultimately will lead to compression of the aquifer, or units above it, and subsidence of the surface. Often this can be severe enough to cause an actual hole to open to the surface. Such a hole poses a threat to structures and human life as well, as being an entry point for contaminants into the water table. Such events already are occurring in the Triassic deposits.

Further information concerning the wells in the Triassic can be found in Table 6 and Appendices A and B.

#### CHAPTER IV

#### GROUNDWATER QUALITY

The use of groundwater for human consumption or industrial use is affected directly by its quality. This quality is determined by naturally-occurring circumstances, but can be altered, both to the good and bad, by man's activities.

### Constituents in Natural Groundwater

As water falls through the atmosphere and percolates through soil and rock material, gases and mineral matter become dissolved and/or suspended in the water. The solids and gases contained in water influence its chemical and physical properties, such as hardness, acidity/alkalinity, corrosiveness, taste, color, odor, etc. Some of the dissolved constituents, such as fluoride (if not excessive), are beneficial; others, such as iron are troublesome. Consultation with the Virginia Department of Health is needed for determining ground-water potability.

The most common constituents of natural groundwater include calcium, magnesium, sodium, bicarbonate, sulfate, and chloride ions.

Other constituents commonly present, but usually in lesser amounts, include iron, manganese, silica, potassium, fluoride and nitrate ions.

Table 4 lists other constituents and their allowable parts-per-million established by the Virginia Department of Health and the State Water Control Board.

<u>Calcium (Ca) and Magnesium (Mg)</u> These two elements, in the form of cations (positively charged ions) are responsible for most of the hardness and scale-forming properties of water. Water containing

TABLE 4
VIRGINIA GROUNDWATER STANDARDS

# Table 4.1 Inorganic Chemicals

(for human consumption)

Substance	Primary Maximum Contaminant Levels (mg/l)
Arsenic (As) Barium (Ba) Cadmium (Cd) Chromium (Cr) Fluoride (F) Lead (Pb) Mercury (Hg) Nitrate (as N) Selenium (Se) Silver (Ag) Cyanide (CN)	0.05 1.0 0.010 0.05 1.8 # 0.05 0.002 10.0 0.01 0.05 5.0

Substance	Secondary Maximum Contaminant Levels (mg/l)		
Chloride (C1)	250.0		
Copper (Cu)	1.0		
Foaming Agents	0.5		
Hydrogen Sulfide	0.05		
Iron (Fe)	0.3		
Manganese (Mn)	0.05		
Sodium (Na)	No Limits Designated, See Text		
Sulfate (SO <sub>4</sub> )	250.0		
Zinc (Zn)	5.0		
Total Organic Carbon	10.0		

# Note. For artificially fluoridated waterworks the minimum concentration of fluoride should be 0.8 mg/l and the maximum should be 1.0 mg/l. The optimum control limit is 0.9 mg/l.

# TABLE 4 cont.

Table 4.2 Organic Chemical

Substance	Primary Maximum Contaminant Levels (mg/l)
Chlorinated Hydrocarbon Insecticides Endrin Lindane Methoxychlor Toxaphene Aldrin/Dieldrin Chlrodane DDT Heptachlor Kepone Mirex	0.0002 0.004 0.1 0.005 0.003 0.01 0.001 0.001 none none
Chlorophenoxy Herbicides 2, 4-Dichlorophenoxyacetic Acid (2, 4-D) 2, 4, 5-Trichlorophenoxypropionic Acid (2, 4, 5-TP or Silvex)	0.1 0.01
Petroleum Hydrocarbons	1.0
Phenols	0.001

Table 4.3 Physical Quality

<u>Parameter</u>	<u>Concentration</u>
Color	15 Color Units
Odor	3 Threshold
рН	6.5-8.5
Total Dissolved Solids (TDS)	500 mg/l
Turbidity	l Turbidity
Alkalinity	10-200

less than 75 milligrams per liter (mg/l) (correlative with parts per million) is thought to be soft and above 75 as hard (Gass, 1978). A great body of literature suggests that there is a relationship between the hardness of water and the incidence of cardiovascular disease. That is, that the softer the water the higher the occurrence. It is indicated that high-hardness waters tend to be low in various harmful constituents: i.e., cadmium, lead, copper, and zinc, which may cause other diseases. All the facts are not complete in either of the two theories, (Gass, 1978). It is known, however, that a high concentration of magnesium can have a laxative effect on the user. High concentrations of calcium, on the other hand, have not been found to be at all harmful.

Sodium (Na) and Potassium (K) In high parts per million these cations can give water a salty taste in combination with chloride. If a person is on a low-sodium diet for hypertension, the water should contain less than 20 mg/l. A sodium-restricted diet should contain less than 2000 mg/day. Normally groundwater is far less than this, rarely exceeding 250 mg/l. It is estimated, however, that up to 40% of all Americans would benefit from a reduced sodium intake; but the quantity derived from water is small.

Iron (Fe) and Manganese (Mn) These constituents cause stains on laundry, cooking utensils, and porcelain fixtures. They also contribute to hardness and may impart an objectionable taste and color to food and beverages.

Bicarbonate  $(HCO_3)$  Bicarbonate ions are responsible for most of the alkalinity in water and combine with calcium and magnesium to form boiler scales.

Sulfate  $(SO_4)$  Sulfate combines with calcium to form boiler scales and may impart a bitter taste to water when present in excess of 500 mg/l. Water containing 1000 mg/l may be strongly laxative in effect.

Chloride (C1) In excess of 100 mg/l, chloride imparts a salty
taste.

<u>Fluoride</u> (F) Concentrations of fluoride of approximately 1 mg/l, depending on temperature, are generally thought to be beneficial by reducing the occurrence of cavities in teeth. Several harmful effects have been attributed to high fluoride water, such as skeletal fluorosis (bone densification), mottling of tooth enamel, mongolism, cancer and birth defects. No evidence of any of these problems, except the first two, is supportive.

Nitrate  $(NO_3)$  All sources of combined nitrogen represent a potential source of nitrate. There are two health hazards related to the consumption of water high in nitrates. These are the possibilities of methoglobinemia in infants (blue babies) and the possibility of the formation in the body of nitrosomines, which may be carcinogenic (Gass, 1978). Studies have shown that infants are suspectible to as low concentrations of nitrates as 10 mg/l.

Specific Conductivity A measure of the ability of water to conduct an electric current. Specific conductivity is an indicator of the relative amount of dissolved minerals in water.

Hardness Other than the discussion of hardness under the Calcium and Magnesium section, it should also be mentioned that hardness causes excessive use of soap. By Virginia standards water with a hardness less than 60 mg/l is considered soft; 60-120 mg/l, moderately hard;

121 to 180 mg/1, hard; and greater than 180 mg/1, very hard.

pH A measure of the hydrogen ion concentration, pH indicates whether water will act as a weak acid or an alkaline solution. The pH scale ranges from 0 to 14 with 7.0 being neutral; greater than 7.0 indicating an alkaline solution, and less than 7.0 an acidic solution. Corrosiveness of water generally increases with a decreasing pH, although highly alkaline water may be corrosive. Corrosive water may attack steel casings or copper plumbing.

As essential nutritive substances, both major and minor (trace) elements are often found in well water. Twenty-five elements are supposedly essential to life (Keller, 1978). At least 14 of the 25 are essential for humans as trace elements. These are Fluorine (F), Silicon (Si), Vanadium (V), Chromium (Cr), Managanese (Mn), Iron (Fe), Cobalt (Co), Nickel (Ni), Copper (Cu), Zinc (Zn), Selenium (Se), Molybdenum (Mo), Tin (Sn), and Iodine (I). These elements, though needed, easily can be too concentrated and then become toxic to humans. The known allowable concentrations of the various elements are presented below where not discussed previously:

<u>Silicon</u> Other than being bound up chemically in soil and rock, thus causing cloudiness, it is not known to be harmful in drinking water.

<u>Vanadium</u> Commonly found in many soils, vanadium appears to be useful in preventing and breaking up excess cholesterol in humans. Excess quantities, however, in combination with other elements, can be dangerous. Apparently concentrations of 0.03 to 0.22 mg/l are safe (McKee and Wolf, 1963).

Copper Metallic copper is unsoluble in water; however there exists many copper salts. Under special conditions the concentration can go as high as 21 mg/l. Dangerous concentrations of copper for human consumption generally are impossible to consume, as the taste becomes very disagreeable. In the body, trace quantities aid in nutrition and the metabolism of iron in the blood. The daily intake of copper is generally 2-5 mg/day. There is no evidence that poisoning has ever occurred as a result of consumption of copper in water (McKee and Wolf, 1963).

Zinc The normal human intake is 10-15 mg/day and is useful in nutrition. High concentrations (variable) are blamed for nausea and fainting in reported cases of occurrence in groundwater. Normally zinc will either be noticed as a greasy coating or by taste before it becomes dangerous.

Selenium Proof of human injury by selenium is scanty, and definite symptoms of selenium poisoning have not been proven; but it is widely believed that selenium is highly toxic to man (McKee and Wolf, 1963). Yet it is essential for nutrition, and recent findings by Dr. Joel D. Wallach, in Chicago, have indicated that a selenium deficiency during pregnancy can cause cystic fibrosis. The poisoning effects are similar to that of arsenic, with the material retained in the liver and kidney. The element also seems to cause a high incidence of cavities in permanent teeth.

Molybdenum The Environmental Protection Agency reports that it is not considered to be a serious pollutant, but it is a biologically active metal. It does serve an important function in the life

processes of phytoplankton (Arnon and Wessel, 1953) but excessive exposure in mammals may interfere with vital, internal, chemical reactions (Dick and Ball, 1945).

<u>Tin</u> Does not occur naturally in groundwaters. Therefore its presence in a well would be indicative of contamination.

<u>Iodine</u> This element helps assure proper functioning of the thyroid gland and the subsequent reduction of goiter. There are recorded cases of human consumption of iodine-rich water at concentrations as high as 19.2 mg/day.

# Water Quality in Loudoun County

Groundwater quality in Loudoun County is variable, but not surprisingly so when one considers the varied rock types encountered in the county. As a general rule water from acidic (granite-like) rocks is "soft" chemically; although it can be corrosive and sometimes iron-bearing. Other high-level, quartz rocks, i.e., sandstones and quartzites, can have similar traits. On the other hand, the basic rocks, basalts, gabbros and such, tend to be high in ferro-magnesian minerals, that is iron and magnesium. Groundwater from the Triassic Basin is generally hard to very hard, very high in dissolved minerals, and tends to be alkaline.

The reader must keep in mind that wells sampled by the State Water Control Board for groundwater quality data vary in age and state of repair. Naturally, older wells, or those recent ones not built to modern standards, are susceptable to contamination from outside sources. These sources can include septic tank drainfields, faulty underground storage tanks and animal lots. This possibility for local contamination may be one of the reasons to account for differences in water quality

within the same rock type.

Several lists showing water quality in the county are included in this report. Table 5 lists the maximum and minimum concentrations and the values for the constituents in the groundwater of the various aquifers in the county, based on chemical analysis available for this report. Appendix B is a summary of water quality analyses of samples taken from individual wells throughout the County. The hardness trends of the groundwater are illustrated in Plate 7 based on the data from Appendix B.

Triassic Lowlands Plate 7 shows the distribution of water hardness in the Triassic Basin. Although a few wells in the western part of this area yield "hard" water, the bulk, of the water is "very hard" as we would expect from the geology. Levels for dissolved solids are high, but in no cases excessive. Also much of the area contains alkaline groundwater, but only slightly so.

Sulphate levels are lower in this area than other parts of the county, and sodium levels are higher. This is, perhaps, due to the sedimentary history of the Basin. Nitrate levels appear to be significantly high in this area, but the reader is cautioned to consider the narrowness of the statistical base and the possibility of local contamination before postulating any theories for this condition. Additionally, it must be remembered that this average is well within acceptable tolerances.

Layered Clastics These rocks also conform to geologic expectations.

Water from these formations tends to be slightly acidic. Although
the samples exhibited all degrees of hardness, it would appear that,

for the most part, water in these rocks is only moderately hard.

Levels for total dissolved solids in the layered clastic rock area

tend to be lower than levels in other areas.

Table 5 indicates that iron levels are higher than in other areas. This may be attributed to the lower pH. Bicarbonate levels are lower in the water in this region, whereas nitrates, chloride, sodium and potassium are consistent with those of the rest of the county.

<u>Layered Volcanics</u> Unfortunately, very few of the wells sampled for groundwater quality were in areas underlain by these formations. For this reason, drawing conclusions from the data presented is a risky proposition.

The data tentatively show that these formations produce water that is slightly acidic. This water appears to be slightly harder than that of the layered clastics, but still tends to be only moderately hard.

Averages for sulphate may tend to be higher, whereas averages for chlorides may tend to be lower than the averages for these compounds in the rest of the county. Nitrate, potassium, sodium, and iron levels appear to be about the same as in those other areas.

Basement Complex These rocks yield water that is slightly alkaline. This is surprising, since we would expect the water from granitic rocks like these to be acidic. It could be that the wells sampled were in close conjunction to mafic intrusions. This relationship would skew the data toward the basic side. Most samples fall into the hard or moderately-hard categories.

Levels for total dissolved solids are in keeping with those in

		•

the rest of the county and are not intolerable. Sulphate, carbonate and chloride levels are high in comparison to the other groups of formations discussed in this report. Nitrate levels are acceptable, with occasional high levels. These high levels may represent localized contamination.

# Groundwater Contamination

Groundwater Contamination refers to the introduction of an organic or inorganic material, foreign to the native groundwater, that tends to limit the useability of the groundwater. If contaminants are added to the groundwater to the extent that it becomes nonpotable or unusable, the groundwater is said to be polluted. Groundwater pollution usually is more serious than surface water pollution, since reversal is very difficult and costly. The self-cleansing mechanisms available to streams, i.e., fresh-water flushing and oxygenation, operate very slowly in groundwater, if at all.

Sources of groundwater pollution are many and varied. Miller and Scalf (1974) presented a synopsis of the subject in a nationwide investigation sponsored by the U. S. Environmental Protection Agency. In Loudoun County, the more significant sources of groundwater pollution include: septic-tank systems, sanitary landfills, sewage lagoons, petroleum spills, leaking pipelines, leaking gasoline storage tanks, improperly-constructed water wells, certain agricultural activities (fertilizers, pesticides, feedlot and barnyard wastes), highway deicing salts and infiltration of poor quality surface water from lakes and streams.

Septic tank drainfields, by sheer volume of waste water discharged, have the greatest potential for contaminating groundwater.

A drainfield is considered failing if the sewage backs up in the house or floods the land surface. A properly operating drainfield acts as a source for groundwater recharge. Miller and Scalf (1974) state that "Studies have shown that in many housing developments recycling of liquid wastes is an inevitable fact of life."

Septic-tank systems commonly are believed to purify sewage.

Actually, the degree of treatment that waste water receives in a septic system is rarely high. In suitable soils, most bacteria are effectively removed by filtration and biological degradation. However, many soils have a low capacity for the purification of waste. Some soil profiles are too thin to assimilate all of the wastes. Moreover, not all of the undesirable constituents of domestic sewage are completely biodegradable. Considering the seriousness of an improper drainfield, and its potential for causing groundwater pollution, land always should be checked by a soil scientist or Health Department official prior to the installation of a drainfield. Miller et al (1974) states, "Under normal conditions of soil pH, efficient removal of phosphates can take place, but chlorides, nitrates, sulfates, and bicarbonates can enter and move freely within a groundwater body."

Nitrates are the most significant constituents (other than bacteria or viruses) of domestic sewage entering groundwater and lowering its quality. In Loudoun County, the nitrate content of natural groundwater is generally low. High concentrations of nitrate most likely indicate contamination. The most likely sources are septic tank effluent and nitrogen fertilizers. In areas where infiltration of nitrogen fertilizers is insignificant, the nitrate content is a good tracer for septic tank contamination. The Triassic lowlands are

probably the most vulnerable area to nitrate contamination in Loudoun County. The soils are relatively thin, and their capability for assimilating and filtering wastes is limited. The shale and sandstone bedrock is highly fractured. The fractures are inclined 15° to 20° from the horizontal, providing a natural conduit to the water table for liquid contaminants, with little opportunity for filtration and quality improvement. The age of a septic system is a significant factor in groundwater quality, since continuous nitrogen loading causes a build-up of nitrate over a period of years. Lot size and housing density are also very important.

Pollution of groundwater by petroleum hydrocarbons is a very serious, common, and unfortunately, growing problem. Most instances are confined to local areas and involve only one or a few potable water wells. However, there have been cases of large-scale groundwater pollution from petroleum, (McKee et al, 1972 and Howard, 1978). In Loudoun County, minor cases have occurred, but, as yet, no major ones. The principal causes of hydrocarbon contamination include: leakage from storage tanks, leakage from buried pipelines, spills or leaks at bulk storage areas, and transportation spills. Many types of petroleum products are involved, but gasoline and heating oil are the most common. Slow leaks can go unnoticed at a service station until wells become polluted, neighbors smell hydrocarbon fumes in their houses, or customers get water in the gasoline they purchase. In very serious cases, such as the village of Port Loring in Ontario, Canada (Howard, 1978) the water coming from household taps can actually ignite.

The clean-up of subsurface petroleum pooling is difficult, costly, and often ineffective. Alternate water supplies often have to be found when a well is contaminated. A publication which offers solutions to problems involving hydrocarbon contamination of ground-water has been prepared by the American Petroleum Institute (1972).

The human threshold for the detection of gasoline can be as low as 0.005 mg/l, so a small amount of gasoline can render a water well unusable. Recent studies have shown that certain bacteria will degrade gasoline in the groundwater environment, if adequate nutrients and dissolved oxygen are available. However, there are cases in which gasoline has remained in soil for more than 70 years (McKee et al, 1972).

It is beyond the scope of this report to provide a detailed discussion of all the causes and effects of groundwater pollution. Groundwater conditions are so variable and potential contaminants are so numerous that most cases are unique. Any known or suspected case should be reported to the appropriate County or State health officials and/or to the State Water Control Board.

TABLE 5

SUMMARY OF ANALYSES OF GROUNDWATER (expressed as mg/l)

Aquifer	Triassic	Layered	Layered	Basement
	Lowlands	Clastics	Volcanics	Complex
NO.3	0.01	0.05	1.3	0.0
	92.1	14.4	9.0	25.0
	5.6	2.6	3.8	3.1
ഥ	0.0 0.32 0.13	0.10	0.13	0.0 0.3 0.19
[2]	1.6 46.0 14.0	1.0 41.0 12.1	0.1 9.0 5.5	1.0
\$0 <mark>4</mark>	0.0	8.2	12.8	0.9
	20.0	15.0	36.5	43.9
	4.3	12.4	24.7	14.7
HCO <sub>3</sub>	11	12	i   i	118
	224	45	i	188
	151	29	i	136
<b>∠</b>	0.1 15.8 1.5	0.1	6.9 0.8 2.7	1.0
Na	0.6	1.0	5.0	1.8
	46.0	50.0	10.4	50.0
	15.2	10.9	8.1	12.6
Mg	0.6 54.5 17.4	30.0 8.9	2.7 18.3 11.4	0.2 32.0 9.5
Ca	1.2	4	9.0	4.8
	636.3	70	61.0	96.0
	55.5	24	30.2	32.0
اعا	0.01 1.1 0.14	0.01 9.8 1.1	0.3 0.3	0.00 3.20 0.56
Hard	52.2	15	33.6	21
	636.3	298	218.2	332
	200.0	98	122.1	121
TDS	17	46	165	42
	364	90	234	360
	175	68	200	184
Spec*	112	27	95	54
	1300	568	440	462
	402	218	243	233
Hd	6.2 8.0 7.1	5.5	6.2 7.3 6.8	6.0 8.1 7.2
	Minimum	Minimum	Minimum	Minimum
	Maximum	Maximum	Maximum	Maximum
	Average	Average	Average	Average

\* Specific Conductivity

				٠
				-
				•
				•

#### CHAPTER V

#### GROUNDWATER DEVELOPMENT

### General Setting

Groundwater is developed widely in Loudoun County in domestic and public water systems. Water wells in the County are of three types: drilled, bored and hand-dug. Most drilled wells are now constructed by rotary drilling machines. Bored wells are constructed by large diameter, (610 to 1220 millimeters (24-48 inches)), boring machines or earth augers. Hand-dug wells, once common in the Piedmont, now are rarely used for potable water supplies.

Drilled wells are usually 101.6 to 457.2 millimeters (4-18 inches) in diameter; the most common size now favored is 152.4 millimeters (6 inches). Steel casing is installed from the surface through the weathered zone and seated into the bedrock. The zone between the outside of the casing and the well bore (annular space) is grouted with cement from some predetermined depth to the surface. This cement grout is to prevent surface contaminants from seeping down the well bore.

Bored wells are common in the Piedmont and Blue Ridge Provinces. In all cases, they are terminated in the weathered zone, since boring machines cannot penetrate hard rock. Bored wells are lined with concrete pipe instead of steel casing. Water enters a bored well at the bottom, which is open. Bored wells are not usually grouted except at the surface, and therefore, are more susceptible to contamination than properly-constructed drilled wells. Bored and hand-dug wells also are more vulnerable to drying up in times of

prolonged drought, since their depths are limited.

### Public Supplies

The construction of public water-supply wells is regulated by the Virginia Department of Health (1974). A well is designated a public supply if it serves 15 connections or 25 people. Public water-wells are required to be cased and grouted to a depth of at least 50 feet (sometimes greater depths are required) and to provide a supply of at least 1.9 lpm (0.5 gpm) per connection.

Public water-supply wells are relatively numerous in Loudoun County. The highest concentrations are in the Leesburg and Hamilton area. The Leesburg wells, of greatest productivity, are located in the Triassic-age limestone conglomerates and diversified units of the Bull Run Formation to the South. The Hamilton wells, on the other hand, are located in the Marshall granite-gneiss and nearby contact zone with the Swift Run formation. By State Water Control Board records, the Town of Leesburg has, or still owns, twenty-five (25) wells. Most of these have been abandoned or capped-off due to becoming dry, or because of inadequate water quantity when they were drilled. Hamilton is credited with owning, or having drilled, sixteen (16) wells, some for subdivisions. At least six (6) of these have been abandoned. Foxcroft School near Leithtown has had seven (7) wells, four (4) of which are operational at present, though rapidly declining. The drillers logs for this area exhibit a very confusing assembly, but it appears that some wells are in granite-like material and some in the Swift Run. There are indications of a fault structure near Foxcroft School as well. Other known, public-well systems are at Lovettsville, Middleburg, and Hillsboro. Round Hill is contemplating

a well-reservoir system.

In addition to these mentioned systems, some subdivisions, notably Moorcones, Potomac Farms, Andover Meadows, and Richland Acres, own their own well(s). Several churches, schools, and light industries are also listed as owning wells. No heavy industry is known to be using wells in Loudoun County.

Considering all the wells recorded in the name of Leesburg, 1.73 mgd output is calculated. The output of other well systems in the County are computed as follows: Hamilton - 0.4 mgd; Middleburg - 0.2 mgd; Lovettsville - 0.35 mgd; and Foxcroft School - 0.13 mgd.

Key Water Wells

Appendix A lists those wells in Loudoun County for which adequate records exist to allow inclusion in the state monitoring-sampling program. The State Water Control Board number is repeated in Appendix B, where applicable. It is immediately noticeable that only 80 wells have been monitored, only 115 are displayed on Plate 8 and only 375 wells are shown in Appendix B. Yet the U.S. Census Bureau estimates that there about 5000 wells in Loudoun County. Unfortunately many wells were constructed prior to the implementation of the Groundwater Act of 1973, discussed later, and to date have not been found. Even when found, a large percentage of owners know nothing about their wells. Also, unfortunately, many modern well owners are not aware of their responsibilities to report, or require the drillers to report, their wells.

Therefore, the availability map and discussions of present and potential water resources are based on a limited statistical base and informed scientific conjecture. Persons using this document are

urged to contact the local office of the State Water Control Board to supply further information about any incompletely-reported wells in Appendix B, or to report the existence of wells not shown. In addition, if a user notices a well in Appendix B which is known to be abandoned, the Board would like to know about it in order to keep its records updated.

### The Groundwater Act of 1973

The General Assembly of Virginia passed the Groundwater Act of 1973 (Chapter 3.4, Title 62.1, Code of Virginia, 1950, as amended)
"...in order to conserve, protect and beneficially utilize the groundwater of this State and to ensure the preservation of the public welfare, safety and health....Concern for adequate groundwater supplies in the future and protection of groundwater quality prompted this legislation."

The Groundwater Act provides for the declaration of a Ground Water Management Area (GWMA) in a geographically-defined region which experiences, or can be expected to experience, problems with interference of wells, water level declines, overdraft of groundwater, or quality of groundwater.

Requests of a declaration of a GWMA can be made by local government, the State Water Control Board, or the public at large. In either case, a detailed study of the proposed GWMA would be made, and public hearings would be convened to determine the necessity and desirability of designating an area a GWMA.

Industrial and commercial groundwater users in a GWMA who withdraw more than 50,000 gpd are required to file a Registration Statement (Form GW-4) with the State Water Control Board. Prior to

1 r		• ,		t s	

the construction of a new well in a GWMA, the owner is required to submit an application and a Permit to Construct a Water Well and to Use Groundwater (Form GW-3).

In <u>all</u> groundwater areas, users are required to submit a Water Well Completion Report (Form GW-2) for new wells. All public and industrial users are required to submit a Groundwater Pumpage and Use Report (Form GW-6) to the Board quarterly. When an industrial or public well is to be abandoned, the owner must submit an Application and Report - Abandonment of Water Well (Form GW-5).

Further explanations are included in the <u>Guide for Water Well</u>

<u>Contractors and Groundwater Users</u> and in the <u>Rules of the Board and</u>

<u>Standards for Water Wells</u>, both publications by the Virginia State

Water Control Board (1974).

# Groundwater Availability

The availability of groundwater in areas of the County is illustrated on Plate 9. This map is based on existing well records, the occurrence of the geologic formations, and their general water-bearing characteristics. Additionally, local areas which are known to be particularly good, bad or suspect are designated.

The map on Plate 9 is general in nature and all unique areas of either high or low productivity cannot be shown. This is particularly true along the major fault systems and in the broad Marshall Formation to the west and the central Triassic Basin to the east. Normally wells in the Marshall zone are poor to fair in yield; although anomalies have been observed in the vicinities of Purcellville, Hamilton, and Wheatland which yielded water in the "good" range. Another high-yield anomaly was noted in the area about one mile south of Sunny Bank. One

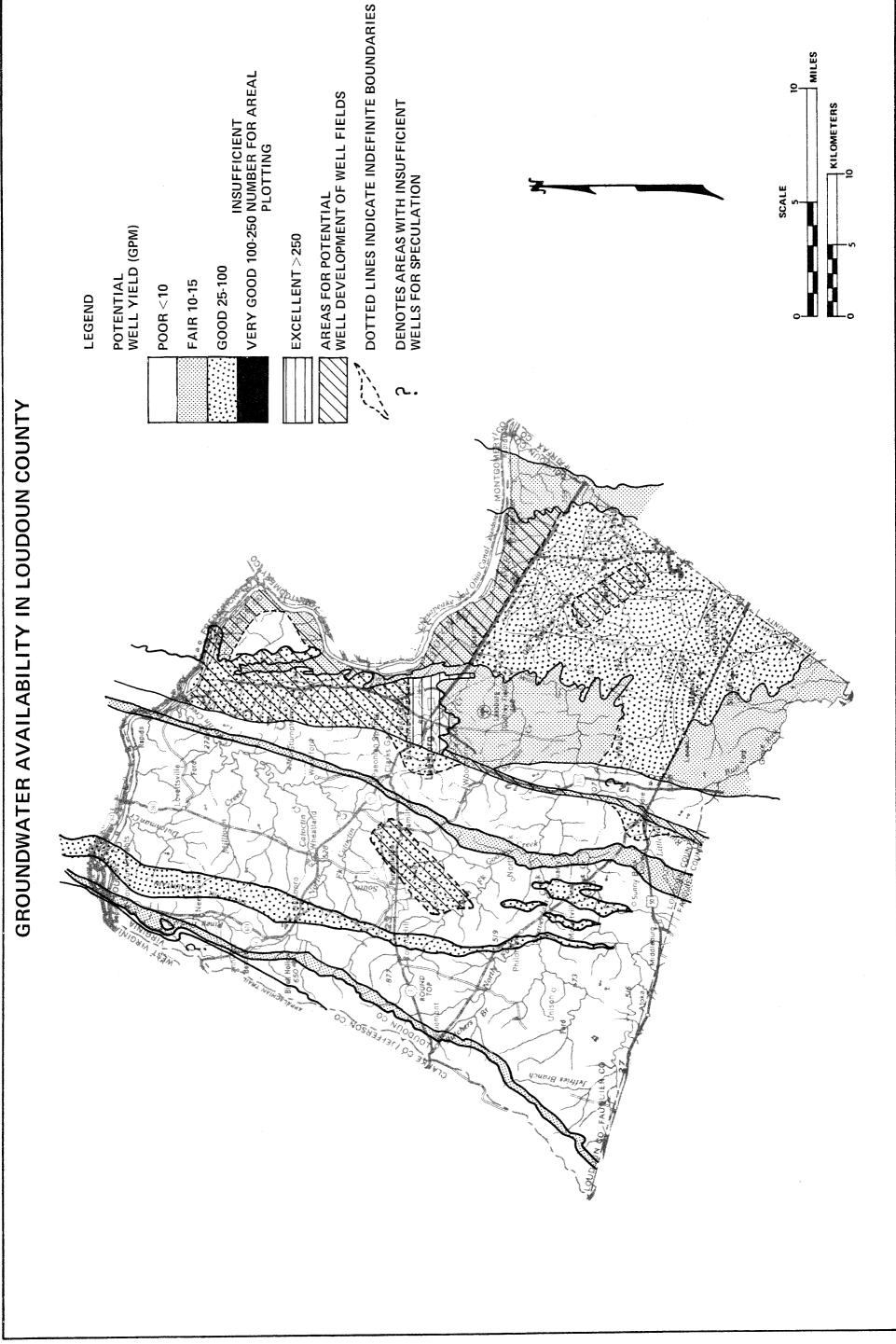
other high-yield anomaly was in the area of the headwaters of Dutchmans Creek, two miles west of Lovettsville. All these above-normal regional yields were in the 95-378 lpm (25-100 gpm) range.

The easternmost Catoctin belt also has two zones of unexpected yield, as shown on the map. These areas lie just west of Aldie and Leesburg. The exact boundaries of these "lobes" are unknown; therefore the configuration shown on the map is speculative.

In addition to the above, there exists the hint of high-yield anomaly all along the Swift Run-Catoctin Contact. It is on this contact that the unidentified, but previously-called "Everona limestone" is found. The width of this outcrop belt is too small to be shown on the plate.

Continuing eastward, the entire Bull Run Fault zone is a potentially-high, groundwater-yielding area and deserves far more investigation and exploitation. The geologic characteristics, i.e., fault dips, fracturing, bedding dips, and proximity to the Potomac River all indicate good groundwater conditions.

Another large area which should receive more attention is bounded by the section of the fault north of Leesburg and Route 7. This area, shown covered with obliquely dipping lines on Plate 9, contains three (3) presently-known groundwater zones. The northern leg of the region, in particular, being Triassic-limestone boulder conglomerate, should be locally highly productive. The shaded area bordering the Potomac River could also be highly productive. This productivity would be due not only to the geology, Manassas Sandstone and Balls Bluff Siltstone, but due to the possibilities of infiltration of water from the Potomac. Public information assistance is requested for all wells in



 •	• 4

this area.

The largest, single water zone in the eastern half of Loudoun is rated "good", with one undefined area containing some wells rated "very good." This smaller area does not, however, contain enough wells to be accurately defined. It should also be realized that there are local areas within the large area that are completely devoid of water. Property owners with small lots should keep this in mind when planning water systems.

Reviewing the data presented on Plate 9, it should be noted that two (2) areas of the map are left blank and marked with a question mark. The first zone is a very narrow slice along the western border marked as £w on the geology map. Physically this area is bounded by the base of the Blue Ridge and the County line and runs south, narrowing to just south of Purcell Knob. The State Water Control Board has no recorded well in this area. Therefore, no speculations other than theory can be made. Based on theory there should be potential for some fair wells in this zone.

The second area is found between the Bull Run fault line and the first basalt flow (most eastern), as defined by K. Y. Lee (1977). The geology in this area is very complex, containing highly-variable, clastic rock units interrupted, or locally covered, by hydrothermally altered basalt flows. At present there exists an inadequate number of wells in this area on which to base any conclusions. The few wells reported have been poor to fair in yield.

As a general statement, it could be said that the water resources are at present greater in the Triassic Basin section of Loudoun County, than anywhere else. If investigated, the Bull Run Fault zone may

exceed the water yields of the Triassic Basin on a local level. Even in a case where a well does not meet a homeowners expectations, there are actions which can be taken in order to assure a dependable water supply. The extracted article from the <u>Water Well Journal</u> of November 1978 (Appendix C) explains three possibilities.

# Well Depths

Considering both the old data reported by Cady (1938) and the modern data assembled by the State Water Control Board, the average productive ranges of wells found in the various geologic units in Loudoun County are given in Table 5. The older wells tend to be much shallower than those drilled with present technology. Also only reported productive wells were used in computing the figures presented. Doubtless there are wells in all the geologic areas which exceed the parameters but have never been reported. It will remain the task of the future to upgrade the data when the base improves.

The data for the Loudoun and Everona Formations were adjusted to compensate for the increasingly-popular and necessary practice of "subdrilling". This is a process of drilling, often as much as 300 feet, below the water-bearing horizon. This provides a storage capacity of 18.2 lpm (1.5 gallons per foot).

It should also be mentioned that the depth of the deepest well shown for the granite unit is highly suspect. One would not expect to find water in granite below 400 feet. In fact, LeGrand in a 1954 report of 520 wells in crystalline rocks showed that the average yields per foot decreased rapidly below 60 meters (200 feet). This appears to be true in Loudoun County as well.

Another item to remember in using the depth vs. geologic unit

guidelines is that in the thinner formations, i.e., Loudoun, Everona (?), Swift Run, Weverton, Harpers, and basaltic flows, a well may penetrate to the underlying unit. In some cases such an event is instantly recognized, but in other instances a driller may not notice the transition.

TABLE 6
WELL DEPTH SUMMARY

Geologic Unit	Depth Range	Average Depth	90% Probability Range
Marshall & all other granite-like rocks	4.5-304m (15-999 ft)	56m (184 ft)	10.7-152.4m (35-500 ft)
Catoctin formation	4.5-215m (15-704 ft)	37.5 (123 ft)*	9.5-102m (31-335 ft)
Loudoun formation <sup>+</sup>	12.8-61m (42-200 ft)	26.8 (88 ft)	15.2-58.5m (50-192 ft)
Everona formation <sup>+</sup>	7.6-36.6m (25-120 ft)	18.6 (61 ft)	9-35m (30-115 ft)
Swift Run formation	16.1-186m (53-610 ft)	70.1m (230 ft)	21.3-152.4m (70-500 ft)
Weverton formation	12.2-157.3m (40-516 ft)	33.5m (110 ft)*	19.5-150m (64-491 ft)
Harpers formation	50-164.6m (165-540 ft)	86.9m (285 ft)	56.4-158.5m (185-520 ft)
Frederick Limestone <sup>†</sup>	11-56.1m (36-184 ft)	insufficient data b	pase
Newark Gp-Leesburg Limestone	8.5-221m (28-725 ft)	54m (177 ft)	10-152.4m (33-500 ft)
Manassas Sandstone	15.2-91.4m (50-300 ft)	37.5m (123 ft)	15.9-62.5m (52-205 ft)
Basalt-flow/bearing Clastics member	18.9-103.6m (62-340 ft)	58m (190 ft)	24.4-103.6m (30-340 ft)
Mixed Clastics	6.1-244.4m (20-802 ft)	44.2m (145 ft)	12.8-163m (42-535 ft)
Balls Bluff Siltstone	6.7-146m (22-479 ft)	35.4m (116 ft)	12.8-99m (42-325 ft)
Triassic diabase (All Areas)	6.7-61m (22-200 ft)	19.2 (63 ft)	9.1-57.9m (30-190 ft)

<sup>\*</sup> Figure adjusted in order to partially compensate for sub-drilling performed to produce a reservoir

<sup>+</sup> After Cady, 1938

#### CHAPTER VI

### CONCLUSIONS AND RECOMMENDATIONS

As originally stated, this report is of a preliminary nature, although it may be used in groundwater resource planning and the development of public and private water systems. It is hoped that in the future, with continued reporting and geologic study, the report can be updated and made more detailed. It should be emphasized that water source plans for specific areas should be supplemented by detailed studies by geotechnical personnel and a plan of test wells. Although several thousand dollars in drilling costs for an inadequate supply of water may seem to be a waste, it is still far less expensive than the development costs of surface supplies.

This report has attempted to briefly explain the geologic diversity, history and geohydrologic environment of Loudoun County. It was stated that the Fault Line and Triassic Basin are the best, potential, water resource areas. The igneous rocks should be avoided, as they are usually dry-to-low producers. The diabase intrusions and feeder dikes are too numerous to be shown on maps of the scale used in the report, and the reader is referred to more detailed maps assembled by U. S. Geological Survey personnel (mentioned in the bibliography).

The relatively-high mineral content of water from the Triassic area may be a limiting factor for industrial use. However, treatment and filtration could correct any problems. Although iron does not appear to be a water problem in any part of the county, the content is lowest in the Triassic sediments. Hardness content is also highest

in the Triassic area, though erratic here as elsewhere. Sulfate content appears to be high in the Middleburg area.

The greatest threat to water quality is activity by man. Pollution and overpumping lead to degradation of both quality and quantity. Such items as septic tanks, drainfields, buried fuel and chemical tanks all have a potential for failure. This is particularly true in high-acid soils as are found in Loudoun County. However, even an innocent and beneficial activity such as weed or pesticide spraying can have an adverse effect on groundwater quality. This is because any water-borne contaminant will seek the lowest, gravity-controlled level, and in a porous and permeable medium this means the groundwater table.

In particular, the uncontrolled spread of septic-tank-using-sub-divisions constitutes the greatest present source of groundwater contamination (Geraghty and Miller, 1978). This practice needs to be more closely monitored and possibly restricted. Potential pollution sources, other than those already mentioned, include underground mining, refuse piles and pits, and slurry ponds. Animal feed lots must also be considered.

Unlike with surface waters, there is no easy, or inexpensive, way to clean groundwater. Pollution, regardless of type, normally will stay until it is no longer dangerous, or until filters and treatment facilities are installed at the withdrawal point. Water has been in the ground a long time. In fact, some of the water sampled at the Dulles Airport well by the U. S. Geological Survey personnel (Rubin, personal communication) was age-dated as approximately 16,000

years old. If the water can stay in the ground this long, so could a water-borne contaminant. Can anyone imagine waiting this long, in order to wait-out a contaminant?

The large number of wells in Loudoun County creates a sizeable management problem, as many wells are possessed by second and third generation owners. Oftentimes these owners have no idea of the characteristics of the well on their property. Such information can be as critical to the useability of the property as the house itself; yet few people even think about it. Modern drilling technology has produced better and safer wells than what was possible a generation ago; but a well-informed owner is a critical link to managing and defending a vital resource. It is recommended that an accurate water well completion report be included with the sales documents for all new homes and that old properties be inspected in order to determine the characteristics of existing wells.

## APPENDIX A

SUMMARY OF GROUNDWATER QUALITY ANALYSES FOR LOUDOUN COUNTY

The computer printout on the following pages lists basic, ground-water quality data available for many of the wells listed in the water well data summary (Appendix B). There are some quality analyses listed for wells not included in Appendix B; however well data is available for these wells and may be obtained by contacting the State Water Control Board's Northern Regional Office in Alexandria or the Headquarters Office in Richmond.

		*
		•

## POOR QUALITY

ORIGINAL(S) FOLLOW

## THIS IS THE BEST COPY AVAILABLE

VCE
DOCUMENT
CONVERSION

## 3. A C 3. 1081800 2 - - - - -4 - 4 - 5 - 5 - 5

## w 51001 AND FIELD SUBVEILLANCE 10 E V S

POCK INGHAM COUNTY SUMMADY OF GRUININGHATEH WALTTY ANALYSES FOR

FROM ANALYSES OF GROUNDWATER DUALITY DATA SUMMARIZES RASIC DATA ORTAINED FROM ANALYSES OF GROUNDWATER. COLLECTED FROM WELLS AND SPRINGS, WHICH ARE ON PERMANENT FILE IN THE OFFICES OF THE VIRGINIA STATE WATER CONTROL BOARD. ADDITIONAL GROUNDWATER ONIALITY INFORMATION FOW MANY OF THESE WILLS AND SPRINGS IS AVAILABLE AND CAN BE OBTAINED BY CONTACTING THE APPROPRIATE PEGIOVAL OFFICE OF THE BUREAU OF SURVEILLANCE AND FIELD STIDIES AT THE AGENCY HEADOUARTERS IN RICHHOND.

## esservers fixplanation of papametens sessesses

SWCH NOT STATE WATED CONTROL ROADD NUMBER - A SEQUENTIAL NUMBEDING SYSTEM USED WITHIN A COUNTY! WHEN REFERRING TO A SPECIFIC WELL USE THIS NUMBER

#ELL. <u>5</u>0 OWNER AND/OH PLACF: IUENTIFIES OHIGINAL OF CURRENT WELL OWNER AND/OF LOCATION

DATE SAMPI. BATE SAAPLED - MONTH AND YEAR IN WHICH WATER SAMPLE WAS COLLECTED.

HYDROGEN ION CONCENTRATION - MASED ON A SCALE OF I THROUGH 14. WATER WITH A PH GREATER THAN 7.0 IS CONSIDERED TO RE BASIC ON ALKALING THE LARGER THE PH VALUE. THE WORE ALKALINE THE WATER. WATER VITH A PH LESS THAN 7.0 IS CONSIDERED TO HE ACIDICS THE SMALLER THE PH VALUE. THE MORE ACIDIC THE SMALLER THE PH VALUE. THE MORE ACIDIC THE WATER. ij

SPECIFIC COMBUCTIVITY - AN INDICATOR OF THE MFLATIVE AMOUNT OF DISSOLVED MINERALS IN MATERI-MIGHEW VALUES INDICATE GHEATER AMPUNTS OF DISSOLVED MINEMALSI UNIT OF MEASUREMENT IS MICROMMO SPEC COND:

T-DIS SOLIDI TOTAL DISSOLVED SOLIDS - INDICATES TOTAL AMOUNT OF DISSOLVED HINEPALS IN WATER! UNIT OF MEASUREMENT IS MILLIGRAMS PFH LITFH

TOTAL MARDNESS - CAUSED BY THE PRESENCE OF CALCIUM. MAGNESIUM. IRON. ZINC. . AND OTHER TPACE METALS. UNIT OF MEACURF IS MILLIGHAMS PER LITEP. CALCIUM-MADDNESS CAUSED BY CALCIUM-MADDNESS CAUSED BY CALCIUM AND MAGNESIUM. WITH A RECENTABLY MESONSIBLE FOR ALMOST ALL MARDNESS. IN JATEM. UNIT OF WEASHIPE IS MILLIGRAMS PER LITER. HAPONESS TOTAL : HARDNESS CA.MG:

THE AMDING OF HARDNESS IN WATER WILL AFFECT THE APILITY OF SOAP TO LATHER OF CLEANSE RECA Of THE TENDENCY OF THE IONS CAUSING HAPDNESS TO WEACT WITH SOAP, THE HIGHER THE HARDNESS WATER, THE MORE DIFFICULT IT IS FOR GOAP TO LATHER.

TOTAL HANDAESS IS GENERALLY DETERMINED BY CHEMICAL TITRATION WHEREAS CALCIUM-MAGNESIUM HANDNESS IS GENEMALLY DETERMINED BY MATHEMATICAL CALCULATION FROM CHEMICALLY-DETERMINED VALUES FOW CALCIUM AND MAGNESIUM, BECAUSE OF THIS DIFFERENCE IN DETERMINATION, THE CALCIUM-MAGNESIUM HARDNESS VALUES FOW SOME ANALYSES WILL BE LARGER THAN THE TOTAL HARDNESS VALUE. NOTE:

PAGAMETERS LISTED BELOW ARE MEASURED IN MILLIGRAMS PER LITER \*\*\*\*\*\*

1 1	FF: IRON	:: Z	MN: MANGANESE	CA! CALCTUM	5
: :	MIL PAGNESTIM	 4 2	WA: SONTUM	KI POTASSIUM	HOI:
¥C97	ACOSE ALCAMANATE	: 70%	SILFATE	כרו כארים10	106

UCON STITABLE (AS NO.)

VIRGINIA STATE WATER CONTROL BOARD BUREAU OF SURVEILLANCE AND FIELD STUDIES

SUMMARY OF GRUUNDWAIER QUALITY ANALYSES FOR LOUDOUN COUNTY

		: ;		2	S0L10	TOTAL CASHG		: :		.!	<u> </u>	<b>:</b>		5.	3	2
28 7	TOWN OF LOVETISVILLE	4 713	7.4		120 123	<b>~ 9</b> ÷	71 0.02 62 1.40	20 0 Z	11.2	25 80 CJ :- \$1	6.1	Ω. Θ.	73	8.0	3.0	0.0
35	TOWN OF LOVETISVALLE	£ 9	<b>3</b>	396	5/2	591 291		Z 0-11	•	_	11.3	0.8	143	54.4	18.5	0.0
1 4 0 0 T T	40 MILLSBORD WATER BEPT 40 MILLSBORD WATER BEPT	7 73 3 71	7°5	ų. 4,	4 U		20 0.00 18 0.08	00.00		0 8	2.1	1.4	10	9.1	4 K	0.4
45	42 NOTHE DAME ACAUEMY #2	69 1	<b>1.</b> 6			7:	0.25		n				66		2.0	0.0
1 94	46 TOWN OF ALUIE	12. 9	9		58		5 0.32	20.0 2		7.0	1.2	1.1	Φ.	4.4	2.0	1.8
1 64	49 LEESHURG	11 <del>5</del>	9:		563	567	0.01	1 0.02	64.1		7.3	3.2	130	12.4	11.5	31.0
56 F	SE EAKTH PRODUCTS	\$ .	F	555		114	4 0.10	ب	26.0	·	40.0	0.3	107		7.0	5,3
n is	S7 DRESDEN FARM	5 74	1.4	96T		124	4 0.10	•	30.0	12.0	16.0	4.0	0.6		7.0	5.8
1 85	SH DRESDEN FARM	\$2. S	1.1	446		261	2 0.10	•	54.0	•	50.0	4.2	108		21.0	18.6
1 65	P K PEARSON	5. 5.	⊃ · ₽	124		243	3 0.10	•	48.0	•	14.0	1.0	201		5.0	
0.9	E M MALKER	5 74	j•9	146		99	6 0.10	0	15.0		8.3	8.0			6.0	10.6
M 19	W L. JAMES	\$.	<b>v</b>	12		41	4 0.10	•	4.0		4.0	0.5	•		2.0	3,5
62 H	H BEKNHOUSE	\$. 5.	<u>د.</u>	271		168	8 0.10	9	46.0	13.0	5.1	1.0	133		3.0	10.6
63	G W PLETCHER	2 C	o.8	317		183	3 0.10	0	0.64	-	4.2	6.0			3.0	12.4
0.49	INOM 0 +9	4/ 9	9.5	'n		33	3 0.10	0	0.6		0.6	6.9	52		0.0	5.8
	65 K JAMES	\$2.0	<b>9</b>	150		T.	1 0.10	9	17.0		0.6	7.8	09		4.0	3.5
<b>3</b> 99	66 W K HARUY	5 74	<b>3</b> ,0	153		1	72 0.10	9	20.0	5.6	10.0	3.8	<b>64</b> .		5.0	12.8
67 5	67 JAMES CUNAKU	\$ .	٠, و	P61		1	01.0 67	9	18.0		11.0	3.9	25		11.0	63.8
89	68 E V LARINHULT	4. d	0.7	411		23	7 0.20	9	13.0		3.0	2.3	54		1.0	3.5
n 69	69 UNISON CHURCH	51.	j•9	462		213	3 j.00	9	61.0	_	25.3	5.8	125		82.0	6.2
70 K	70 K H BLAIR	42 /	7.0	151		50	85 0.10	9	25.0		7.5	1.5	99		4.0	15.9
71 H	71 H & JENKINS	1 74	0.0	148		88	8 1.70	9	23.0		10.5	0.4	J.		•	a

NUTE--ALL ZEROS (UU.00) - ANALYSED, NUT DETECTEUT ALL NINES (99.99) - COULD NUT BE STORED, REFER TO ANALYSIS

STUDIES B O A R C SURVEILLANCE AND FIELD CONTROL WATEX VIRGINIA STATE **-**8 U K F A U

LOUDUUN COUNTY SUMMARY OF GROUNDWAILH QUALITY ANALYSES FOR 

SWCH. GWNER AND/OR PLACE NO	DATE SAMP	ž.	SPEC	1-015 50L10	HARDNESS TUTAL CA+MG	Ţ	Z E	V U:	9±	ď Z	×	ALK	804	ರ	NOO
72 DH DIAMANT	1 74	9	30		51	0.20		15.0	3.5	8.8	2.4	56		5.0	18.2
73 C W MCINIOSH	47. 0	Ĭ°Į	029		379	01.0		78.0	45.0	15.5	2.1	310		17.0	21.3
74 C S PEAHSON	9	7.5	450		599	0.10		0.06	18.2	18.3	1.7	272	4	14.0	3.5
75 w beleu	<b>5</b> 2 9	7.3	410		522	0.10		36.0	33.0	37.5	2.5	192		29.0	15.5
76 H J MEYLE	4/ 9	7.5	0 44		887	0.10		63.0	32.0	18.0	2.0	238		30.0	
77 J B ELLISON	72 9	9	375		189	0.10		46.0	18.2	17.3	15.8	135		16.0	
78 J A MONELAND	71 1	1.6	210		156	02.0		40.0	13.7	0.6	4.5	119		12.0	1.3
79 J A ATWELL	1 74	6.7	961		119	0.10		39.0	5.5	9.	5.4	85		0.9	4.0
80 C ATHEY	1 74	8.1	270		172	0.10		43.0	6.51	8.9	1.0	129		13.0	20.4
81 H SHUTWELL	7 1	7.9	580		340	0.10		105.0	2.61	12.8	S • 0	554		16.0	20.8
82 M COLBERT	1 14	7.0	552		96	2.10		23.0	91	21.5	5.2	110		16.0	0.0
83 A WILLIAMS	7. 0	7.5	360		161	0.10		0.99	0.	13.3	9.0	144		10.0	46.5
84 K FOX	47 B	4:	560		151	0.10		52.0	3.	6.3	٠. د	128		5.0	13.7
BS H P HANEY	1 74	7.3	1300		635	0.10			54.5	46.5	1.3	117		12.0	0.4
86 BUSH	1 74	7.3	720		426	0.10		115.0	34.0	24.5	1.1	273		7.0	1.3
87 J R HARRISON	\$1.1	7.6	200		282	0.10			19.0	29.3	S•0	258 258		20.0	13.7
88 J FOUCHE	1 74	7.3	280		336	01.0		73.0	37.5	23.8	8.0	233		26.0	22.6
89 A ANDERSON	1 74	1.4	450		265	0.10		17.0	17.8	15.0	1.0	180		19.0	18.6
90 HOBERT THOMPSON	10 74	P.	372		232	01.0		71.0	13.4	14.7	0.1	212		7.0	I. 3
91 ALPREU L PHILLIPS	52 1	0°.	125	95	80	1.10		24.0	8	2.6	8.0	11	ō•0	3.0	6.0
92 HAKKNESS	1 75	<b>9</b> :	iis	7.	7.2	01.0		21.0	8.	2.0	8.0	99	0.0	2.0	0.0
93 LYLE C HARTHAN	1 75	<b>D</b>	140	<b>401</b>	80	0.10		22.0	<b>6</b> •0	4.7	1.1	99	0.0	0.6	0.0
94 M G WHITE	52 1	j•į	220	154	131	0.10		38.0	10.3	4.2	8.0	129	4.0	5.0	11.5

NUTE--ALL ZEROS (UU.80) - ANALYSEU, NUT DETECTEU; ALL NINES (99.99) - COULD NUT BE STOKED, REFER TO ANALYSIS

V LHGINIA STAȚE WAȚEK ÇONȚROL BOAKD HUREAU OF SURVEILLANCE AND FIELD SŢUDIES SUMMARY UL GROUNDWAIER QUALITY ANALYSES LOR LOUDOUN COUNTY

NOTE--ALL ZEROS (00.00) - ANALYSED, NUT DEJECTED; ALL NINES (99.99) - COULD NOT BE STOKED, REFER TO ANALYSIS

VIKGINIA STATE WAFER CONTROL BOARU BUREAU UP SURVEILLANÇE AND FIELD STUDIES SUMMARY OF GROUNDWAFER GUALITY ANALYSES FUR LOUDOUN COUNTY

SWCB OWNER AND/ON PLACE	SAMP	ž	SPEC	i-uis souiu	HAKUNESS TUJAL ÇA•MG		∓ π	MN	₩.	ď Z	*	ALK	804	5	E 00
166 CHARLES HUMMER	3 78	7.6	052		15	120 0.20	20.0 0:	2 32.0	0 10 0	11.0	3.0			0.6	
171 PLEASANT VALLEY UMC	8/ 8	j•/	330		77	149		37.0	0 14.0	11.0	1.0	139		12.0	
174 CARROLL MILBOURNE	2 78				417		67.0	<b>3</b>							
175 SAMAH BUXLEY	2 78				104										
177 LEE WILLIAMS	4 78	<b>0</b>			2 <u>8</u> 2							237	0.6	26.0	14.2
182 FOXCHUFT SCHOOL	8/ 4	7.3	320		7	136 3.20	0.00	45.0	0.0	7.0	3.0	84		2.0	0.0
188 WALTER H MCDERMIFT	9) 5	7.0	639		ř	304 0.00	10 0.01	1 76.0	0 48.0	18.0	0.0	237		46.0	8.0
189 L B JAMES	8 <i>1</i> . c	7.4	513		72	246 0.00	00.00	0 74.0	0 15.0	12.0	0.0	223		16.0	0.0
190 CAMEY	8/ s	7.1	571		72	262 0.70	00.000	0.65 0	0.82 0	16.0	1.0	223		25.0	3,1
229 DAVID BOWEN	9/ 9	7.4	044		Ñ	218 0.00	00.0 0	0 61.0	0 16.0	7.0	0.45	162		0.0	17.7
231 EVERETT W LOWE	9/ 9	1.9	661		•	85 0.30	00.0 0	0 21.0	0 · 8 · 0	5.0	1.0	44		n•6	13.3
247 S W MUNKOE	1 78	<b>6</b>			•	45		10.0	0.5.0	1.0	3.0	. 51		1.0	
248 JUE BENNET	1 /8	4.0				36 0.10	01	8.0	0.4.0	2.0	0.2	41		1.0	
265 DOUGLAS MYEHS	8/	7.6			Ā	101 1.90	*0 · 0 · 0 ·	0.62 %		21.0	9.0	011 0			5.3
266 ALEXANDER SWENEY	82.	<b>3</b> ,			87	331 0.90	20 0 0¢	6،08 در	0 32.0	0.05	3.0	184			61.7
270 N F CUPELAND	92.	B. 0				05.0 77	10.0 00	21.0	0.0	0 - 2 - 0	3.0	080		2.0	
272 DONALD I VINTS	1 78	7.3	181		2.	06.0 77	90 0°12	12 23.0	0 5.0	0.9 0	0 5.0	25 0		3.0	
273 KICHAKD ROUEKS	8/ /	5.7	/RT		0 9	86 0.10	10 0.04	04 28.0	0.4 0	0-8 0	0 5.0	0.6 0		1.0	0.0
275 GEURGE SHANNON	R/ /	6 3				92		19.0	0 7 0	0 12.0	3.0	0 25		20.0	
276 WILLIAM C BENDEK	8/ /	6.0			146	651		0.44	0 12.0	0 19.0	0 2.0	0 25		18.0	
279 EUWIN CHEEL	8//	D .	115		N	280		0.84	0 27.0	0 15.0	0	526	•	20.0	
281 JUSEPH PRENDERGAST	8) )	9	107			25	0.05	05 11.0	0.0	0 8 0	0 3.0	71 0		11.0	
285 JAMES KAYLUR	8)	7.0	307			152 9.	9.80 U.S.	54 43.0	0 11.0	0.6 0	0 2 0	0 108	~	41.0	

NUTE--ALL ZERUS (UU.UU) - ANALYSEU, NUT DETECTEU; ALL NINES 199,99) - COULD NOT BE STORED, REFER TO ANALYSIS

BOAKU CONTROL BUREAU UP SURVEILLANCE ANU \* A T IT K A I N I O N I V

STUDIES

FIELU

ಗ 13.0 0.9 6.0 5.0 204 ALK 302 **7** 39 9 1.0 2.0 2.0 0.6 13.0 7.0 18.0 **4** LOUDUUN COUNTY 7.0 42.0 5.0 æ 4.0 110.0 20.0 0.96 20.0 18.0 S SUMMARY UF GRUUNDWAIER QUALITY ANALYSES FOR 0.02 ĭ 0.10 0.02 330 0.80 0.04 0.30 Ŧ, 78 HARDNESS TUTAL CA:MG 364 7.0 9 336 301 9 56 1-015. 501.10 SPEC 215 299 175 213 865 7.6 1.1 ... ٧. ĭ UA IE SAMP 8 7 8 8/8 87 8 8/ R 82 × SWCB OWNER AND/OR PLACE NO 373 H W ALEXANDER 386 GOLDEN'S WELL 363 CLATTERBUCK 375 WILLIAMS 374 DANIELS

EON

NOTE--ALL ZEROS (00.80) - ANALYSEU, NUT DETECTED; ALL NINES (99.99) - COULD NOT BE STOKED, REFER TO ANALYSIS

107.0

7.0

36.0

22.0

## APPENDIX B

## SUMMARY OF WATER WELL DATA FOR LOUDOUN COUNTY

The computer printout on the following pages lists basic well data for wells in Loudoun County. This printout is updated frequently to include information from new Water Well Completion Reports and old well records. The former is being submitted by water well drillers and the latter is found by State Water Control Board personnel. The information under the heading "Aquifer" may be cross-referenced with Table 2, Chapter III.

			•
			•
			•
			٠
			•
			٠

## z Ľ ⋖ لغا 0 Σ L. 9 9 ⋖ 0 Z α Σ Œ z 0 0 ں œ -z ¥ ند **⊢** ၁ ပ ¥ ¥. A ⊣ E TATE S 0 ⋖ I 2 I 9 2 I >

u.

C K FI D

## LOUBOUN COUNTY

SUMMARY OF WATER WELL DATA FOR

THE FOLLOWING LIST OF WELL DATA SUMMARIZES BASIC DATA OBTAINED FROM WATER WELL COMPLETION REPORTS WHICH ARE ON PERHANENT FILE IN THE OFFICES OF THE VIRGINIA STATE WATER CONTING BOARD. AUDITIONAL INFORMATION FOR MANY OF THE WELLS IS AVAILABLE AND CONTROL BOARD. AUDITACITUS THE APPROPRIATE REGIONAL OFFICE ON THE BUREAU OF WATER CONTROL MANAGEMENT AT THE AGENCY HEADGUARTERS IN RICHMOND.

## \*\*\*\*\*\*\*\* EXPLANATION OF PARAMETERS \*\*\*\*\*\*\*\*\*

STATE WATER CONTROL BOARD NUMBER - A SEQUENTIAL NUMBERING SYSTEM USED WITHIN A COUNTY; WHEN REFERRING TO A SPECIFIC WELL USE THIS NUMBER SWCB NO:

WELL P OWNER AND/OR PLACE: IDENTIFIE'S ORIGINAL OR CURRENT WELL OWNER AND/OR LOCATION

YEAR COMP! YEAR IN WHICH WELL CONSTRUCTION WAS COMPLETED

GEOLOGIC/GAMMA GEUPHYSICAL. LOG: TYPE OF LOG ON FILE FOR WELL! U = DRILLERS, E = ELECTRIC, G = L = INTERVIEW LOCATION FORM, M = WOALITY MONITORING WELL. ELEVATION - MEASURED IN FEET ABOVE MEAN SEA LEVEL

TOTAL DEPTH: TOTAL DEPTH DRILLED, IN FEET, WITH RESPECT TO LAND SURFACE

BEDROCK: DEPTH TO BEDROCK, IN FEET, WITH RESPECT TO LAND SURFACE

MAXIMUM AND MINIMUM DIAMETER OF CASING, IN INCHES, USED IN WELL CASING DEVEL ZONE: DEVELOPED ZONE - INTERVALS, IN FEET, WHERE AQUIFERS TAPPED AND/OR SCREENED

CONSISTENT WITH TABLE 2 AND GENERALIZED GEOLOGIC MAP OF LOUDON COUNTY IN TEXT WATER-BEARING UNIT! ABBREVIATIONS USED INDICATE GEOLOGIC AGE OF UNIT AND ARE AGUIFERS

STATIC LEVEL: DEPTH, IN FEET, TO WATER WITH RESPECT TO LAND SURFACE! MEASUREMENTS TAKEN WHEN WEND WELL IS NOT PUMPED AND ARE GENERALLY THOSE RECORDED ON COMPLETION DATE

YIELD: REPORTED OR MEASURED PRODUCTION, IN GALLONS PER MINUTE

DIFFERENCE, IN FEET, BETWEEN STATIC LEVEL AND PUMPING LEVEL! I.E., REPORTED MEASURED DROP, IN FEET, IN WATER LEVEL DUE TO PUMPING DRAWDOWN:

SPECIFIC CAPACITY - YIELD PER UNIT OF DHAWDOWN EXPRESSED AS GALLONS PER MINUTE PER FOOT OF URANDOWN SPEC CAPAC:

- DURATION OF PUMP TEST, IN HOURS, FROM WHICH THE PRECEDING THREE ITEMS WERE DERIVED IKS.

USE OF WATER OR WELL UNDER CONSIDERATION! DOM = DOMESTIC, PUB = PUBLIC, GOV = GOVERNMENT, IND = INDUSTRIAL, COM = COMMERCIAL, INS = INSTITUTIONAL, ABD = ABANDONED, DST = DESTROYED, IRR = IRRIGATION, RCH = ARTIFICIAL RECHARGE USE:

VIRGINIA STATE WATER CONTROL BOARD BUREAU OF WATER CONTROL MANAGEMENT

SUMMARY OF WATER WELL DATA FOR LOUDOUN COUNTY

	USE	PUB	₩ 0 0 0
	HRS S	4 1 1 5 5 4 4 8 8 8 8 4 4 1 1 5 5 1 1 5 5 5 5 5 5 5 5 5 5 5 5	
	SPEC	27.88 .01 .03 .04 .04 .05 .05 .05 .07 .07 .07 .07 .07 .07	
*	DOWN	168 168 423 64 40 1115 103 40 1115 103 103 103 103 103 103 103 103 103 103	
*	YIELD	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	20 100
***	STATIC	2850 2850 2850 2850 2850 2850 2850 2850	8 4
· · · · · · · · · · · · · · · · · · ·	AGUIFER	TREE STATES OF THE STATES OF T	E E
* * *	ZONE TO	64 29 29 83 150 150 150 151 161 161 175	
*	DEVEL	51 52 90 90 120 150 150 150 150 150 150 170	
* * * * * * * *	CASING MAX MIN	<b>~ w e w o a a a a a a a a a a a a a a a a a a</b>	ហហ
**	BED- ROCK	2 4 4 5 8 8 8 7 9 11 9 9 8 8 8 4 4 5 8 8 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9	
**	TOTAL DEPTH	285 285 285 285 285 285 285 285 285 285	300
	ELEV	6 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	5.38 5.00
*	907	000000000000000000000000000000000000000	نہ نہ
***	YEAR	6 6 6 6 6 8 8 8 9 8 9 8 9 8 9 9 9 9 9 9	99 90
******	OWNER AND/OR PLACE	USGS-MYERS #1 USGS-MYERS #2 USGS-MYERS #3 USGS-MYERS #3 USGS-MYERS #3 USGS-MYERS #3 CRIDER & SHOCKEY H B HARRIS HARRY MCDERMITH DUNCAN H READ DUNCAN H READ DUNCAN H READ T G SLATTER & SON INC A T & T G SLATTER & SON INC A T & T	U LYNN CORNWELL #1
	SWCB	- 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4	1 1

VIRGINIA STATE WATER CONTROL BOARD BUREAU OF WATER CONTROL MANAGEMENT

## SUMMARY OF WATER WELL DATA FOR LOUDOUN COUNTY

USE	
HRS	ru 4
SPEC	19.20
D C C C C C C C C C C C C C C C C C C C	25 9 7 1
YIELD	
STATIC	90 0 1 2 3 3 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
AQUIFER	PC C C C C C C C C C C C C C C C C C C
ZONE	29.5
DEVEL FROM	293
CASING MAX MIN	N 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
BED-	35 20 20 10 10
TOTAL	390 248 248 400 190 1100 100 100 100 100 100 100 100
ELEV	2000 2000
P00	33131313131313131313131313131313131313
YEAR	588
OWNER AND/OR PLACE	J LYNN CORNWELL #3 CLAUDE CHUKCH TOWN OF HAMILTON #6 LEESBURG-PAXTON WILLIAM BREHM TOWN OF HAMILTON #7 JOHN ROBISON J WILLSON J WILLIAMS R I JENNINS DENEN R BLAIF R I JENNINS DK. DIAMANI C W MCINTOSH C S PEARSON W BELEU J A MORELLIAMS R FOX H P HANEY UN HARRISON J R HARRISON
SWCB	4 4 4 4 4 4 W W W W W W W W W W W W W W

VIKGINIA STATE WATER CONTROL BOARD BUREAU OF WATER CONTROL MANAGEMENT

UMMARY OF WATER WELL DATA FOR LOUDOUN COUNTY

USE	
±8S	1
SPEC	6.00 10.00 11.0.00 11.0.00 11.0.00 11.00 10.00 1
DORAW NA	50 50 50 50 50 50 50 50 50 50 50 50 50 5
YIELD	25 25 25 25 25 25 25 25 25 25 25 25 25 2
STATIC	7
AQUIFER	TARES
ZONE	1160 2370 260 260
DEVEL FROM	80 3 4 5 5 0 4 4 0 0
CASING MAX MIN	0000 r 000000000 000000000000000000000
BED- ROCK	882 116 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1
TOTAL DEPTH	125 125 126 127 128 128 128 128 128 128 128 128 128 128
ELEV	40000000000000000000000000000000000000
P00	
YEAR	2
OWNER AND/OR PLACE	A ANDERSON ROBERT THOMSON ALFRED L PHILLIPS HARKNESS LYLE C HATMAN M G WITH BWILLIAM SAVOPOULOS LOUIS V HALL OBED O BROWN DOUGLAS A PERICH J D THOMAS VILLIAM SAVOPOULOS LOUIS V HALL OBED O BROWN DOUGLAS A PERICH J D THOMAS VILLIAM SAVOPOULOS LOUIS V HALL C CARE-POINTERER TOWN OF LEESBURG #1 TOWN OF LEESBURG #2 LEESBURG-PITEMBN WELL LEESBURG-PITEMBN WELKE LEESBURG-PITEMBN WELKE LEESBURG-PITEMBN WELKE LEESBURG-PITEMBN WELKE LEESBURG-PITEMBN WELKINS L MANTINN THOMAS L KENT JOHN WILKINS
S.W.CB NO.N.	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0

z Σ z ⊃

USE

	ģ	Ŷ	80	œ	m	60		2	~		m	m	-					,	m e	m (	2	۳)							٣	-	۰,	า			~	, r	n (*	, ~	י	,	۰ ۱	ı		84	
	( (	CAPAC	• 05	60.	• 03	.02		.11	60.		• 25	.02							:	•13	20.								.01			1.00				F 0 '	י י	3	•	.03		1			
*	i v	* Z O O	550	550	535	615		159	123		235	300							;	215	300								278			001			080	27.5	0.00	000	22	0.50	200	9			
***	2 1 2	Y 1 EL D	15	20	50	5		Ś	12	50	9	o	20		<b>→</b>		9	S :	0	9	•	ဓိ	<b>10</b>	ď	u u	n			ın	i	ກຸ	007	<u>y</u>	0 1	•	<b>u</b> 5	3 -	۰ •	า	a	o u	•		10	
***	, ,	STATIC	30	20	10	35		45	42	9	40	<b>4</b> 3	100					ı	14	į	32	010	S.						47		į	97		,	2 0	ה נו	0 0	) u	<b>4</b> U	9	0 t	מ		τ	
00UNTY	1	AGUIFER	CPCC	CPCC	TRBF	TRBF	₽C.¥	CPCC	TRBF	CPCC	TRBF	TRB	CPCC	PCE	CPCS	P C M	<b>≱</b>	A.	PCG	TRB	PCM	TRBF	TRB	920	0 20	E 0	200	200	PCM	PC	)       	TRLC	200	٦ د د د	200	E (	E 7	E 0	E 1	0000	5000	ו ז ז	200	TREC	
LOUDOUN COUNTY		ZONE TO																																											
DR LO		DEVEL FROM																																											
SUMMARY OF WATER WELL DATA FOR LOUDOUN COUNTY		CASING MAX MIN	•	0	• •	<b>.</b>	•	9	• •	• •	• •	9	9		•	•	•	9	•	<b>.</b> 0	•	:•	<b>6</b> 0			ıs v	o		9			9	<b>0</b> 0	<b>0</b> 0 ·	۰ م	φ.	φ,	۰ م	9	•	۰ م	٥ 4	0	•	,
WATER WELL		BED- ROCK	35	50	, ~	202	i	17	7.0	•	52	25					•	4	15	'n	20	2							25		30	2				20.5	80 :	9	38	Ċ	9 7	ប្ម		28	j
OF WA		TOTAL DEPTH	550	550	7 C	615	9.80	250	145	040	250	340	540		620	345	100	125	250	240	355	160	290	62	90	390	621		595	120	330	160	175		080	270	562	200	235	7	245	230	000	725	,
SUMMARY OF		ELEV	475	4	4 4	325	ر ا ا	400	5.0	1040	410	365	069	200	485	513	420	280	900	340	460	365	330	375	320	250	ν. υ.	217	430	200	310	325	745	745	580	373	084	4 0	365	000	700	1310	200	000 7.5	,
**************************************		507	a	9	· c	o C	3 0	<u>م</u>	<b>a</b>	s C	2	۵	۵	ب	0	۵	٥	٥	٥	۵	۵	٥	_1	. د	. فــ	. د	. د	_ نـ	۵ ا	L	۵	٥	ب	ب	، ب	<b>3</b> i	۰	2	9	، د	۰ د	۵.	. د	ے د	۲
\$ \$ \$ \$ \$		YEAR	11	7.7		7.7	14	9.4	7.0	11	- 82	78	11	9	7.8	1.1	7.7	11	7.7	7.7	78	17	75		09	99			7.8		78	78	99	9	63	7.8	78	78	78	30 y	18	9 5	6	0 00 1 12	-
**************************************		OWNER AND/GR PLACE	TOWN OF LEFABLISHED	, <b>i</b>	, L	OF LEESBOAG	ACE CANADA POR SELECTION OF THE PORT OF TH		STATE OF THE STATE	COCK AND SCORET	CHENTER A KNAK	THOMAS P. JORDAN	L P & ANN BURNETT	HUDDLE RESTAURANT	JIMMY COX GCF34	MARK REED	PAUL ROSS	DAVID HUBERT	LARRY A PATRICK	JOHN RICCA	ED FELLOWS	C M ZACK	T A LYNCH		CHANTILLY C S	J LYNN CORNWELL INC	JOHN CLATTERBUCK	BLUEMONT GEN STORE	F H DEBUTTS	HILLTOP GARDEN CENTER	BILL RAPSON	LEESBURG COUNTRY CLUB									STANLEY E RUSSELL			SARAH BUXLEY	
		S&CB NO	5.5	3 6	† U	n 4	9 6	200	2 0	707	7 7		7 7	144	145	146	147	148	149	150	151	152	153	154	155	156	15		160	161	162	163	164	165	166	167	168	169	170	171	172	173	7/1	175	2

CONTROL MANAGEMENT CONTROL WATER WATER STATE J. VIRGINIA BUREAU

COUNTY
LOUDOUN
FOR
DATA
MELL
WATER
S
SUMMARY

USE		W W W
HRS	u &uwwa anana	~ ~
SPEC	. 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	.04
DORAW	192 470 470 100 120 120 175	385 145
YIELD	10 20 20 20 20 20 20 20 20 20 20 20 20 20	400
STATIC	44 64 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8	35
AQUIFER	C C C C C C C C C C C C C C C C C C C	CPCS TRBB
20NE T0		
DEVEL FROM		
CASING MAX MIN	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	O OO OO
BED-	15. 20 1 1 1 1 1 2 2 0 2 8 2 2 0 2 8 8 2 9 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8	, y ce o
TOTAL DEPTH	200 200 200 200 200 200 200 200 200 200	160
ELEV	$\begin{array}{c} \text{www} \\ \text{u} \\ $	292
r00	000000000000000000000000000000000000000	200
YEAR	77	78
OWNER AND/OR PLACE	LEE WILLIAMS EDWARD B DAWSON JR HUMMER JOR HUMMER MOORCONES SUBD #15&16 MOORCONES SUBD #15&16 MOORCONES SUBD #15&16 MOORCONES SUBD #15&16 MALTER H MCDERMITT LE JAMES CARY RALPH CURTIS V B BORAH LETCHER AL WILL HELENA HERNDON PAINTER FOTH MARK E CHAMBERLAIN WILLIS H MARTIN JR WILLIS H MARTIN JR WILLIS H MARTIN JR WILLIS H REISNER EDWARD A CHOK C J MOORE EDWARD A CHOK EDWARD A CHOK EDWARD A CHOK EDWARD A CHOK ELOUIS & SAKAH HERTZOG SYCOLINE 1ST BAPT CH R J DEHART HERBERT H FARRIS W GREY PRICE HUTCHISON & MASON M N LYON CHARLES MCLEHEE JOHN F MCOLONEY GLENN A TEUKSBURY LEONARD C GOWMAN PIEDMONT LAND CO.	E V FINN (42, PARC 56A) RIDDLE CONST
SWCB	178 178 186 186 186 186 186 187 187 188 188 188 188 188 188 188 188	220

0 VIRGINIA 8 U R E A U

USE	######################################	X X X
HRS	7 A B B B B B B B B B B B B B B B B B B	
SPEC		
DORAW	180 145 125 55	
YIELD	25 8 2 2 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	-
STATIC	20 0 0 5 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6	
AGUIFER	COCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC	O O O O O O O O O
ZONE TO		
DEVEL FROM		
CASING MAX MIN	ላ <sup>3</sup> 440 ወወወወ ወ	36
BED- ROCK 1	29 29 10 20 20	
TOTAL B DEPTH R	205 150 63 63 63 63 63 63 63 63 63 64 60 60 63 60 63 60 63 60 63 60 63 60 63 63 64 60 63 63 64 64 63 64 64 64 64 64 64 64 64 64 64 64 64 64	30 100 500
ELEV	22.00 mm m	395 390 380
907		ددد
YEAR	4 4 4 6 6 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8	38
PLACE	DHRER DY KREITZER JR SH	GONSETH ERS
OWNER AND/OR PLACE		263 JEWELL 264 GEN & MRS GUNSETH 265 DOUGLAS MYERS

USE

SWCB

	× × ×	z ⊃ - •	A 0	4 4 - 3 0	בט בט בע בע	* P C C E	2 O	α ∢ Σ z	ν Γ. Σ Σ	M A D				
		•	SUMMARY	OF WA	3	٥	OR L	<u> </u>	COUNTY					
2000年	* * * * * * * * * * * * *	*	\$ \$	* * * * * *	* * * * * * * * * * * * * * * * * * * *	* * * * *	* * * * *	牵	****	******	********	* * * *		
OWNER AND/OR PLACE	YEAR	٥٥٦ -	ELEV	TOTAL DEPTH	BED- ROCK	CASING MAX MIN	DEVEL FROM	ZONE TO	AQUIFER	STATIC LEVEL	YIELD	DRAW	SPEC	HRS.
	;	•	,	1		į								
ALEXANDER SWENEY L T MCCRACKEN SH	0 0	ب در	370	175		98 4			0 0 0 0 0	ţ	u			
WARREN BOGGESS	}	. نــ ا	024			36			P.C.		1			
AMAKI N FI COPFI AND	9	<b></b>	370 235	125					¥ 20	7	S			
DEANE QUINNEY	523	. فىد ل	535	140		36			ž Ž	-	n'			
DONALD T VIKTS DICHARD DOGEDS	20	۔ ب	200	6		,			E C		;			
GEORGE SHANNON	0 99 - W	ر ر	430	300 85		o o			ž č	- -				
GEORGE SHANNON	58		440	120		• •			PCS	2	•			
WILLIAM C BENDER	37 8	ے ب	510	50 10 10 10 10 10 10 10 10 10 10 10 10 10		<b>8</b> 4	976	776	PCM	ć	C	u		,
HASTENING FARM LTD	78	۵ ۵	552	210	<b>x</b> 0	o∵•o	ĵ.	) )	P.C.M.	25	205	185		ግ ሊ
EDWIN CREEL	45	: بـ	340	135		•			TRO	i			-	ı
UOSEPH PRENDERGAST		د د	9 60			vc			1885		4			
JOSEPH PRENDERGAST	16	<u>ئ</u> ـ ا	305	230		• •			TRBF		52			
JAMES S KAYLOR	,	٠,	320	4 v		84			A C					
CHURCH OF REDEEMER	2 9	ب د	320	ດ <b>ເ</b> ດ		<b>1</b> 0 <u>c</u>			3 3					
AUBREY VIRTS	78	٥	480	140	0+	· •	130	131	PCA	52	15	130	•00	٣
RANDY WATERMAN	80 2	Δ.	810	445	20	<b>,</b>			CPCC	38	4	420		2
ROBERT BELL	9.2	c	432	105	20	o •o			188F		4 0			
GROVER CRANDALL 47F-2	7.8	۵	200	270	S	· <b>v</b> o			CPCC	45	20	255	•19	٣
REX H SHUMAKER	78	٥ د	540	909	į	•			PCM					
	8 2	<b>&gt;</b> =	570	202	ς, γ				S T C		15			
J G SCHOOLER	77	۵	505	200	į	<b>.</b>			E S	30	,			2
F ADAMS LISTEPHEN ETT TAKE		<b></b>	400	190					TRBF					
ALFRED WILLIAMS		اس ا	370						TKBF					
ALFRED WILLIAMS		. ب	365						TRBF					
WALIER ENIERPRISES CECIL E TEATS			400						TRBF					
W B TRITTIPOE		ا ل	350						188					
THOMAS DELASHMUTT	20	. بـ	370	110		9			TRB					
RUDION CANIACOZENE		ـ لـ	395	0		,	5		CPCC		9			
PAUL MCDONALD	- 22	ـ د	001	100 200 200 200		۰ ۵	9		E 1		<b>5</b> -`3			
ž	78	۵ د	270	220	15	<b>.</b> •			18g	38	. S.	192	.07	~
	7.8	2	322	115	18	9			THLC	B ~	5.5	78	.70	٦,
TOWN OF LEESBURG COWS	78 78	22	300	145 115	20 52	<b>6</b> 6	•		TRLC TRLC	98 88 88	325 500	72 38	4.51	4 4 8 4

VIRGINIA STATE WATER CONTROL BOARD BUREAU OF WATER CONTROL MANAGEMENT

SUMMARY OF WATER WELL DATA FOR LOUDOUN COUNTY

USE	
H S	
SPEC	
DRAW	475 1197 1119 1119 1119 1119 1119 1119 11
YIELD	4 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2
STATIC	444W4WW4W WWWW WW WW WW WW WW WW WW WW W
AQUIFER	CPC CAMB B S S S S S S S S S S S S S S S S S S
DEVEL ZONE FROM TO	232 360 100 440 440 255 200 200 200 210 210 210 210 210 210 210
CASING DE	NA N
BED- ROCK	113 81
TOTAL DEPTH	200 200 200 200 200 200 200 200 200 200
ELEV	$\begin{array}{c} \mathbb{U} \\ $
F06	
YEAR	2
OWNER AND/OR PLACE	LEESBURG DIZEREGA JAMES D TEMPLE CONRAD S MALLOY MARVIN A WALKER J CANTRALL JOHN W HILL JR FRANCIS NEELY DEAN RATHBURN L D BURGER B C HAGNOR JOHN C MINOGUE LAROCK UNKNOWN F PAUL GRASSO F JOHN C MINOGUE LAROCK UNKNOWN F PAUL GRASSO CLAYTON DAVIS BETTY L MOYERS JOHN A PARRICK UNKNOWN F SHEHEN CLAYTON DAVIS BETTY L MOYERS JOHN A PARRICA MANTE PILLION WHITE PILLION WHITE PILLION ANDREW COATES SOBETTY L MOYERS JOHN DILLON DOUGLAS LAWSON MADISON HOUSE KENNETH B MARSHALL JOSEPH MCDUROUGH DOUGLAS LAWSON MADISON HOUSE KENNETH B MARSHALL JOSEPH MCDUROUGH DOUGLAS LAWER JONETTE M HUCK THELMA N POORE JOHN DILLON DUGLAS LAWER JAMES W HARRIS
S#CB NO	3110 3110 3110 3110 3110 3110 3110 3110

## SUMMARY OF WATER WELL DATA FOR LOUDOUN COUNTY

ш.	
USE	
HRS	84
SPEC	
DORAN	350 150 185 185 330
YIELD	6 2 3 3 3 3 3 3 3 3 3 3 3 3 3 3 3 3 3 3
STATIC	120 20 30 30 30 30 40 40 40 40 40 40 40 40 40 40 40 40 40
AQUIFER	CCPC CCPC CCPC CCPC CCPC CCPC CCPC CCP
ZONE TO	
DEVEL	1180 235 1130 130 225 215 215
CASING MAX MIN	24 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6
BED-	28 9 8 9 9 9 5 8 8 5 5 8 8 8 9 9 9 9 9 9
TOTAL	1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1
ELEV	80.0 40.0
907	
YEAR	2777777777777888888887778877888778888778888
OWNER AND/OR PLACE	WILLIAM DOEHKEN BRUCE PATTERSON FRANK AMOS HENRY RAMEY R H RANDOLPH GEORGE HORKAN WILLIAM S STOKES III JOHN HANNA DEAN RATHBUCK FITZWATER CHARLES MILLER FRED CALHOUN HALTERBUCK FITZWATER CHARLES MILLER FRED CALHOUN HALTERBUCK FITZWATER CHARLES MILLER HALTERBUCK FITZWATER CHARLES MILLER HALTERBUCK FITZWATER CHARLES MILLER HALTERBUCK FITZWATER HALTERPUCK HACKLEY HALTERS-PIEDMONT MT JOHN W DEAN HWALTERS-PIEDMONT MT JOHN W DEAN HIDEO I HARA BRADLEY B JOURNELL INSTITUTE CORP HIDEO I HARA BILLY B JOURNELL INSTITUTE CORP HIDEO I HARA BILLY B JOURNELL CRAWN ROBERT F KAN EDWARD BURLING EDWARD ENGER JOHN ENGER JOH ENGER JOHN ENGER JOHN ENGER JOHN ENGER JOHN ENGER JOHN ENGER JOH
SO ON	3356 3356 3356 3356 3366 3366 3366 3366

# C IN I B STATE \* ATEX CONTRCL BOARCES CONTRCL BOARCES CONTRCL BOARCES CONTRCL BOARCES CONTRCL BARREN

USE

α α	D C		ы 4	SUMMAR	ARY OF	T A X	т х х <u>х</u>	C O N T	טא מ	UUUUUNN	N A G E	는 건 네 호					
	*****	*****	*****	* * * * * * * * * * * * * * * * * * * *		٠	* * * *	*****		* * * * *	***	************************************	* * * *	*			
OWNER ANDZOR PLACE YEAR LOG ELEV TOTAL COMP DEPTH	106 ELEV	ELEV	ELEV	>			BED- ROCK	CASING MAX MIN	DEVEL	20NE T0	AQUIFER	STATIC LEVEL	YIELD	UCE A SE	SPEC CAPAC	I S	
0 440	0 440	044	044		~~~	20	130	9			SPCC	80	1			p4	
PRESSWAY CORNER 78 D 270	0 270	270	270		17	30	10	9	160		TRN	40	18				
78	00			9	70 V	0 1	5.				FCMA A	35	20	160	•12	→ -	
HIGGINS SECSSPARIS 78 U	<b>3</b> 2			ان خ د	t v	വ	5 0	ه ه			າພູ	ს 4 ა ბ	~,	543 400		,	
THERN LLOYD 78 D	۵			چ ر	202		5	Φ.			PCMA	38	10	180	• 05		
76 D	2			o ·	0 7	0	15	•			PCMA	52	'n	225	• 02		
055 0 87	0 530	5.30	5.30		36	0	Ξ.	<b>Q</b>			PCMA	38	75	338	£0°		
0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	0 p 1	0 p 1		165	n c	2 0	ο ι			O J	0 4 0 4	0 e	0 t 0	35	~ <b>,</b> ~	
0 82	۵ ۵			<u>ان</u> د	3		· ac	0.0			F C W	36	) 30	061	40.	• -	
A 78 D	٥			0	316		14	ف			PCMA	44	30	280	.02	4	
£Y 78 D	a			.C	54	'n	30	đ			PC	<b>.</b> 2	∧ı	200		~4	
L 530	L 530	530	530		91	0:	50	٠ ۵			CPCC	đ	300			<b>~</b> .	
STATES 00 1 245	240	0 0 0 0 0 0 0 0 0	0 0 0 0 0 0 0 0 0		ž á	<b>5</b> 10	0.7	ρχ	100		X X	Ç	v			<b>-</b>	
78 0 435	0 435	4.35	4.35		77		63	۵ ه	200	205	200		30			~	
78 0 435	55.4 6.5 6.5 6.5 6.5 6.5 6.5 6.5 6.5 6.5 6.5	435 35	435 35		<u> </u>	0 (	20 9	۰ م			1 X C						
370	370	370	370		ē	9	o <b>1</b>	٥			T N	1 1					
VS CENTER #2					304	<b>.</b>		ø			CPCC						
HONICAN		045	0.40 0.40	<b>.</b>	7	_		1			Z 2	36	120			9	
2 20		335	335	<u>ت</u> د	400			ه د			Z Z Z	2	7			1	
SAFEWAY #2 78				0				0			TRN						
TOWN OF LEESBURG PL#2 78 D	٥			0 (	Š	<b>.</b>	<b>6</b> 4	۰ م			Z Y Y	7 5	25	202	0.	2	
NOTRE DAME ACADEMY 78 D	S			nc	9	טי. ט	n oc	0 40			F CMA	n so	17	580	.02	7	
NOTRE DAME ACADEMY 79 D	) )			. 0	4.1	ı un	4	• •			PCMA	38	17	390	<b>*</b> 0 <b>*</b>	m	
ALICE JACKSON 78 D	a			ũ	17	7	140	•	175	177	TRI	40	0.7			-	
F WARREN 79 0	<b>3</b>			0	45	ır.	82	٥			သ ပည် ပည်	4 i	<b>†</b>	400	• 01	-	
٥٥	٥٥			٠ و د	9 -	0 0	01	<b>0</b> 1			ပ သ (	15	- 1 - 1				
R HIGH MCCALL FY 78 0 550	0 0 0 0 0 0	0 0 0 0	0 0 0 0		2 🗟	<b>,</b>	10	0 6			r a	200					
000	<b>S</b>			2 9	1 -	o u	2 0	0 4			ر م	, r	3 4	126	131	-	
1NC 79 0 620	000	000	000		1 4	0 មា	<b>7 7</b>	0.0			າວ		15	125	.12	<b>-</b>	
SUE PULE0 79 0 485	0 485	485	485		30	ں:	38	٥			PCMA	43	m	582	.01	-	
267 0 82	964 0	495	495		00	S	15	٥			CPCS	29	cO :	585		7	
0 87	2			õ	20	9	50	9			IXI	40	7				
သ း	သ း			بر ا	36	in a	32	•	250	2910	ပ (၁)	or o	ო -	340		-	
ASHLAND ACKES INV CO 79 D	20			<u>n</u> ~	2,7	<b>.</b> c	υ γ.	o o	יחכ	<b>†</b> 0	4 E 2 E	٥	7 7			-	
JOE LESBURG 78 D	0			0	16	٠ ٥	104	• •			PCMA	50	12				
ADM WILLIAM SMALL 79 D	۵			. 5	16	ı in	17	O			T CMA	27	15	140	• 10	~	

VIRGINIA STATE WATER CONTROL BOARD BUREAU OF WATER CONTROL MANAGEMENT

COUNTY
LOUDOUN
FOR
DATA
- K
WATER
9
SCAMARY

	USE		
	HRS	- 2	
	SPEC		
* * * *	DRAM	200	
***	YIELD	12 4 2 1 1 3 1 1 0 1 1 0 1 1 1 1 1 1 1 1 1 1 1	
***	STATIC	8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8	
****	AQUIFER	TTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTT	
***	ZUNE TO	180 50	
***	DEVEL FROM	30 220 400 178 100	
· 李安安安安安安安安安安安安安安安安安安安安安安安安安安安安安安安安安安安安	CASING MAX MIN	00000000	
* * * *	BED- ROCK	12 52 6 6 7 100 100 100 100	
* * * * * * *	TOTAL DEPTH	530 300 460 340 220 560 300 120	
***	ELEV	0444 0460 0460 0464 0464 0464 0464 0464	
* * * *	۲00	00000000	
	YEAR	70 70 70 70 70 70 70	
*********	OWNER AND/OR PLACE	A C ECHOLS JR ALLEN DEV CO MARK RIFFLE ROBERT D BASSFORD RICHARD O GORE RICHARD TO GORE LENA MORRISON JR LENA MORRISS ROGER ZYLSTRA JAMES BAGALY	
	SWCB	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	

## APPENDIX C

"Using Large Diameters to Increase Well Storage" \*

You don't have to buy bottled water just because your well yields less than one gallon per mintue.

Even a  $\frac{1}{2}$  gpm well can provide a reserve of 720 gallons per day more than enough to supply a family of four for two days.

Some aquifers, such as "tight" shales or limestones, just aren't porous enough to permit prolonged pumping. Especially when local ground water conditions have not been adequately assessed, low yield wells may result. These will require some form of storage.

Storage system design requires an accurate estimate of present and future water needs. Ways to calculate water use requirements are described in "Planning for an Individual Water Supply," and "Water Systems Handbook," (Both available from the National Water Well Association). Light industry relying on low yields can forecast water requirements through process engineering studies.

## Large diameter well storage

The best storage area for a reserve water supply is the well itself. Contractors can increase a wells' storage capacity simply by drilling larger diameters. Although water production capacity does not increase dramatically with well diameter, every added diameter inch radically increases storage capacity. (Table 1)

The table shows that a low yield well with 4-6" diameter could not supply a family of four requiring 200 gpd. Periods of heavy use would quickly dewater the well. An 8-inch well with

75 feet of water-filled casing would just meet this family's needs.

Any well 10 inches in diameter or larger with 50 feet of water

storage would exceed the family's requirements.

Keep in mind that water use is highly variable. In the consumer's best interest, drillers should suggest that casing storage capacities be at least 50 percent greater than apparent requirements.

## Pressure tanks

Water system pressure tanks can also help provide storage. These tanks prevent rapid cycling of the pump motor, and provide additional pressurized water storage. Both of these functions are critical in increasing storage capacity of low yield wells.

An oversized pressure tank can be used to provide supplemental water storage. Tank drawdown is increased to meet the system's demands.

Proven well yield, pump capacity, maximum water requirements and operating pressure and pressure range numbers will determine pressure tank size. Storage capacity is usually only a fraction of total daily water use; and only 10 to 30 percent of the tank is full at a given time.

When peak usage rates exceed the well's capacity, a larger pressure tank can be used. An intermediate storage tank may be more appropriate if storage requirements exceed 700 gallons. In any case, the water system should be able to supply the peak use rate continuously for one hour in domestic systems, and two hours for light industrial and farm applications.

In some cases, both a large diameter well and an oversized

pressure tank may be required. This combination would provide daily domestic needs from the pressure tank, and an emergency storage supply in the well casing.

Above ground non-pressurized tanks

Light industry and agriculture can also take advantage of well storage capacities. Required storage volumes are much larger than domestic applications: they should be able to meet water needs for 24 to 48 hours. Water supply wells must also yield more - 5 to 10 gpm for a 10,000 gallon storage tank is sufficient.

Since systems for light industry and agriculture must also provide water for uses at varying pressures, booster pumps and pressure tanks are used.

Water is pumped into the non-pressurized storage tank by the well pump. Pumping head is only required to overcome pipe friction loss and elevation between the pump and storage tank "full" level.

A float-level is used to control pump operation. The pump "on-off" range is usually 100 gallons. (After the float switch turns the pump off, 100 gallons of water must be drained from the tank before the pump will be reactivated.)

Water levels in a low yield well supplying such a storage system must be closely controlled. Solenoid activated pump switches attached to water level sensors in the well are used for this purpose.

The sensors are located at two levels in the well: the lower sensor a few inches above pump intake, and the upper one just below the well's minimum static water level. When water level drops

past the lower sensor, a circuit is broken, activating the solenoid switch and turning off the pump. When the well recovers its natural static level, the upper sonsor resets the lower sensor circuit and starts the pump.

Control of pump motor speed is another way to prevent dewatering of low yield wells. Pump motors are allowed to pump only 85 percent of the wells' production capacity. Since it cannot pump more water than the well produces, the pump will not dewater the well.

Pump motors for low yield well systems must be continuous duty type, since the pump may be required to operate around the clock.

TABLE 1

Volume of Water Storage In
Water Filled Casing Length

Nominal Casing Size	Inside <u>Diameter</u>	Storage per linear ft.	50' Section	75' Section	100' Section
4" 5" 8" 10" 12" 14" 15"	4.03" 5.05" 6.07" 8.07" 10.19" 12.09" 13.25" 14.25"	0.66 gal 1.04 gal 1.50 gal 2.66 gal 4.24 gal 5.96 gal 7.16 gal 8.29 gal 9.49 gal	33 gal 52 gal 75 gal 133 gal 212 gal 300 gal 358 gal 415 gal 475 gal	49.5 gal 78 gal 113 gal 200 gal 318 gal 447 gal 537 gal 622 gal 712 gal	66 gal 104 gal 150 gal 266 gal 424 gal 600 gal 716 gal 830 gal 949 gal
18" 20"	17.18" 19.18"	12.05 gal 15.01 gal	603 gal 751 gal	904 gal 1126 gal	1206 gal 1501 gal

<sup>\*</sup> Reprinted by permission of Water Well Journal Publishing Company.
Copyright 1978 All Rights Reserved

## GLOSSARY OF TERMS

ACIDIC ROCKS:

A general descriptive term referring to igneous rocks which contain greater than 66 percent silica  $(SiO_2)$ .

ALLUVIUM:

A general term for sediments deposited during recent geologic time by a stream or river.

AMYGDALOIDAL:

Describes volcanic rocks characterized by cavities formed by gas bubbles and subsequently filled with minerals differing from the rock mass.

ANTICLINE:

An arch or fold in rock strata that is convex upward.

AQUICLUDE:

A geologic formation which is not permeable enough to furnish an appreciable supply of water to a well or spring.

AQUIFER:

A geologic formation, group of formations, or part of a formation capable of supplying water to wells and springs in usable quantites. An aquifer is unconfined (water table) if the surface of the water table is at atmospheric pressure or confined (artesian) if the upper surface of the groundwater is under pressure in excess of atmospheric pressure due to the presence of an overlying, confining formation (aquiclude).

ARKOSE:

A granular sedimentary rock composed of large grains of quartz and feldspars from the disintegration of acidic rocks.

ARTESIAN SURFACE: Level to which water will rise in a closed column

due to hydraulic pressure.

AUGEN: Refers to large mineral grains which have the

shape of an eye.

AUGEN GNEISS: A banded metamorphic rock characterized by the

presence of large phenocrysts, "augen", of various

minerals.

BASALT: A dark, ferromagnesian-bearing igneous, extrusive

rock.

BASE FLOW: The flow of groundwater into a stream.

BASIC ROCKS: A general term describing igneous rocks which are

low in silica and relatively high in dark

colored minerals.

BASEMENT: Complex of crystalline igneous and metamorphic

rocks which underlie the sedimentary sequence in

any region and appear to have no other units

below them.

BEDDING PLANE: A plane surface separating layers in stratified

rocks.

BEDROCK: A solid rock exposed at the surface (outcrop)

or overlain by unconsolidated materials.

BIOTITE: A common black mineral of the mica group.

BOULDER: A rock fragment greater than 256mm (10.08 inches)

in diameter.

BRECCIA: A rock composed of highly angular fragments often

caused by crushing or movement along a fault.

CALCITE:

Calcium carbonate, a common mineral which occurs in cracks and cavities in rocks and as the main constituent of limestone.

CAMBRIAN:

A term used to describe a geologic age, 570-500 million years ago, and also those rocks which were formed during that period.

CASING:

Pipe, either metal, plastic, or in some cases cement, which is placed in a drilled hole to prevent the hole caving and to seal out surrounding contaminants.

CATCHMENT:

The area comprising the actual water intake area for aquifer recharge and all areas that contribute surface water to the intake area.

CHLORITE:

A sheet-like dark green mineral belonging to the mica group.

CLASTIC:

A term referring to rocks composed of fragments of pre-existing rocks.

CLAY PAN:

A layer of compacted but not cemented clay which is relatively impervious to water.

COBBLE:

A rock fragment with a diameter of between 64mm  $(2\frac{1}{2} \text{ inches})$  and 256mm (10.08 inches). Often rounded from stream transport.

COLLUVIAL:

Unconsolidated masses of sediments that are gravity deposited.

CONE OF DEPRESSION: The depression in the level of the water table caused by pumping water out of the aquifer. The depression usually takes the form of an inverted cone around the point of withdrawal.

CONFINING LAYER: An impervious layer which restricts water to the

aquifer in which the water is already present.

CONGLOMERATE: A rock composed of rounded fragments of pre-

existing rock cemented together by another

mineral.

COUNTRY ROCK: Refers to the rock into which an igneous (molten)

mass intrudes.

CROSS BEDDING: Lineations or laminations within a layered

sequence which are oblique to the direction of

the main strata.

CRYSTALLINE ROCKS: Rocks whose constituent minerals have a definite

recognizable crystal form.

DIABASE: A fine to medium grained dark colored rock similar

in composition to basalt.

DIASTROPHISM: The earth movement process of mountain building.

DIKE: A tabular mass of igneous rock which cuts through

adjacent rocks. Usually caused by the intrusion

of a magma.

DIP: The angle between the surface of an inclined

stratum and the horizontal plane.

DOLOMITE-MAGNESIUM

CARBONATE:

A common rock forming mineral similar in composition

to limestone but containing magnesium and

sometimes iron.

EPIDOTE: A pistachio-green calcium silicate mineral

commonly found in metamorphic rocks.

EUGEOSYNCLINE: A large area formed as a geosyncline which

contains volcanic material.

FACIES: A distinct unit of sedimentary rock distinguished

from adjacent units by differences in lithology,

fossil content, structure, etc.

FAULT: A plane of fracturing along which movement has

taken place in some direction.

FEEDER DIKE: A smaller offshoot of a dike which shows the

same cross cutting relationship and connects

to the main dike.

FELDSPAR: Important rock forming minerals of the aluminum

silicites.

FELSITE: A term given to an igneous rock consisting of

extremely fine crystals of quartz and feldspars.

Some larger crystals may be present.

FLUVIAL: River-deposited sediments.

FOLD: A curve or bend in rock strata.

FORMATION: An assemblage or rock masses grouped together

into a unit that is convenient for description

or mapping.

FRACTURE: Breaks in rocks due to intense folding or faulting.

FRAGIPAN: A resistant layer of clay with calcareous cement

which is imperivous to water.

GABBRO: A deeply emplaced magmatic rock similar to

granite, though darker.

GARNET: A mineral group which is often indicative of

metamorphism.

GEANTICLINE: A broad uplifted section of the crust of the

earth.

GEOHYDROLOGY: The science that relates geology and the water

beneath the earth.

GEOSYNCLINE: A large syncline.

GNEISS: • A metamorphic rock characterized by alternate

bands of granular and platy or, more generally,

light and dark minerals.

GPD: Gallons per day.

GPM: Gallons per minute.

GRANITE: A rock which forms below ground and consists of

alkalic feldspars, quartz, and minor amounts of

biotite, mica and other feldspars.

GRANULITE GNEISS: A high temperature, banded, metamorphic rock

consisting of even-sized grains and containing

mica and hornblende.

GRAYWACKE: A general term referring to a grayish rock

composed of angular fragments of granular igneous

in a clay-silica groundmass.

GREENSCHIST: Both a degree of metamorphism and a schistose

green rock consisting primarily of chlorite.

GREENSTONE: A general term given to basic igneous rocks, such

as basalt, which contain abundant chlorite,

epidote and hornblende.

GROUNDMASS: The material surrounding the crystals and grains

in a rock.

GROUNDWATER:

Water occurring in the zone of saturation beneath

the land surface.

HYDRAULIC LIFT:

The difference between the top level of an artesian

aguifer and the level to which a column of water

will rise from this aquifer due to hydraulic

pressure.

HYDROLOGY:

The science that relates to the water of the

earth.

HYDROTHERMALLY

ALTERED:

Refers to the processes whereby rocks and their

constituent minerals are changed by exposure to

mineral rich waters associated with magmas.

IGNEOUS:

Rocks or minerals that solidified from molten

rock (magma).

IGNEOUS INTRUSION:

Refers to the forceful emplacement of molten

rock into existing rock.

IMPERMEABLE:

Refers to a substance or strata which will not

allow water to pass through it.

INTRUSIVE:

Refers to igneous rocks which have penetrated into

or between older rocks while molten but have

solidified before reaching the surface.

JOINT:

A fracture in rock along which no appreciable

movement has occurred.

KAOLINITE:

A common aluminum silicate clay. Kaolin, a rock

composed chiefly of kaolinite is used to make

porcelain and china.

LITHOLOGY:

The composition and structure of rock.

LPM:

Liters per minute.

MAFIC:

Refers to ferro-magnesian minerals or, more

generally, dark colored minerals.

MASS WASTING:

Refers to the movement of large rock or earth

masses due to gravity.

METAMORPHIC:

Refers to any rocks derived from pre-existing

rocks in reponse to pronounced changes of

temperature, pressure, and chemical environment.

MICA:

Refers to a group of thin-sheetlike minerals

which are an important constituent of many

metamorphic rocks.

MIOGEOSYNCLINE:

A large geographic area formed from a syncline,

but lacking volcanic materials.

MGD:

Million gallons per day.

MONZONITE:

A rock which contains orthoclase and plagioclase

feldspars in roughly equal amounts. This rock

forms below the ground surface and usually con-

tains small amounts of quartz.

M.Y.B.P.:

Million Years Before Present.

NORMAL FAULT:

A fault in which the upthrown side of the fault

does not overhang the downthrown side.

ORDOVICIAN:

A term used to describe a geologic age, 500-430

M.Y.B.P., and also those rocks which formed

during that period.

OVERBURDEN:

Generally used to refer to all unconsolidated

materials which overlie bedrock.

PALEOZOIC:

A geologic era which extends from 570 M.Y.B.P.

to 225 M.Y.B.P.

PERCOLATION:

Movement of water through the interstices of

rocks or soils except movement through large

openings such as solution channels.

PERMEABILITY:

A measure of the ability of a rock, sediment

or soil to transmit water.

PERVIOUS:

Refers to a material which will allow water to

pass through it.

PHENOCRYST:

A large, distinct crystal in a finer grained

groundmass.

PHYLLITE:

A micaceous metamorphic rock similar to slate,

but lustrous.

POROSITY:

The property of a rock, soil, or other material

of containing spaces or voids; the ratio of the

void space to the total volume in a given sample

of rock or soil, expressed as a percentage.

POTENTIOMETRIC

SURFACE:

The level to which water will rise in a cased

well; describes the water table surface or the

artesian surface depending on the confining

conditions of the aquifer.

PRECAMBRIAN:

Refers to a geologic age, prior to 570 M.Y.B.P.,

and those rocks which formed during that time.

PUBLIC SUPPLY:

As defined by the Virginia Department of Health,

a water system serving 25 individuals or more

than 15 residential connections.

PUMPING LEVEL: Depth to water in a well when the well is being

pumped.

PYROCLASTIC: Rocks formed of volcanic debris, i.e., ash,

cinders, bombs.

QUARTZ: Silicon Dioxide  $(SiO_2)$ . One of the most abundant

rock forming minerals.

QUARTZITE: A metamorphic rock consisting of greater than

95 percent quartz.

RECHARGE: The addition of water to an aquifer by natural

infiltration or artifical means.

RELIEF: The difference in elevation between high and low

points on the earth's surface.

RHYOLITE: This igneous rock forms above ground and consists

of silica and feldspar. Glass content (silica)

is often very high.

RUNOFF: That part of precipitation that appears in streams;

includes surface runoff and groundwater flow

that reaches streams. Groundwater runoff is a

measure of the change in groundwater storage

and indicates rate of groundwater recharge.

SAPROLITE: Refers to weathered, highly decomposed rock.

SCHIST: A metamorphic rock with somewhat-parallel grain-

orientation giving a banded appearance by color

or mineralization.

SEDIMENT: Material borne and deposited by water, wind, or

glaciers.

SEDIMENTARY:

Refers to rocks formed from the consolidation of

sediments.

SERICITE:

Fine grained mica often found in schists; often

appears to be a clay.

SHALE:

A laminated sedimentary rock consisting primarily

of compacted clay size grains.

SILT:

An unconsolidated sediment whose grains range in

size from 1/16mm to 1/256mm in diameter.

SLATE:

A metamorphic rock characterized by distinct

cleavage and fine grains.

STATIC LEVEL:

Depth to water in a well when the well is not

being pumped.

STRUCTURE:

Refers to the physical features present in an

area or in an individual rock.

SYENITE:

An igneous rock which forms below ground and

consists of feldspar and mafic minerals.

SYNCLINE:

A fold in rock strata that is convex downward.

Pertaining to, or designating the external earth

forms resulting from deforming forces such as

compression, tension, etc.

TALC:

A very soft mineral used in talcum powder,

insulators, etc.

TECTONIC:

Refers to the processes and results associated

with deformation of the earth's crust.

TERRACE:

A level or gently inclined surface bordering

a stream which represents a former level of

the stream. Terraces are composed of alluvium

produced by renewed downcutting of the flood plain or valley floor by the stream.

TEXTURE:

Physical aspects of a rock including grain size,

shape, arrangement, etc.

TOPOGRAPHY:

The surficial aspects of an area, i.e., flat,

rolling, hilly, etc.

TRIASSIC:

A term used to describe a geologic age, 195-225

M.Y.B.P., and also those rocks which were formed

during that period.

TUFF:

A rock composed of small particles of volcanic

debris which are welded together.

UNCONSOLIDATED:

A sediment that is loosely arranged or whose

particles are not cemented together.

WATER TABLE:

The upper surface of the zone of rock or soil

saturated with ground water.

WEATHERING:

A general term referring to the chemical and

physical processes whereby rocks are broken

down into smaller particles or basic minerals.

## **BIBLIOGRAPHY**

- American Geological Institute, 1960, "Glossary of Geology and Related Sciences, with Supplement." 2nd Edition, Washington, D.C.
- Arnon, D. I., and Wessel, G., 1953, "Vanadium as an Essential Element for Green Plants." Nature. 172, pp. 1039-1040.
- Bloomer, R. O., and Werner, H. J., 1955, "Geology of the Blue Ridge Region in Central Virginia." Geological Society of America Bulletin, Vol. 66, No. 5, pp. 579-606.
- Brown, W. R., 1954, "Structural Framework and Mineral Resources of the Virginia Piedmont." Virginia Geological Survey Reprint Series No. 16.
- Brown, W. R., 1958, "Geology and Mineral Reources of the Lynchburg Quadrangle, Virginia." Virginia Division of Mineral Resources, Bulletin 74, 99 p.
- Brown, W. R., 1970, "Investigations of the Sedimentary Record in the Piedmont and Blue Ridge of Virginia, in Studies of Appalachian Geology, Central and Southern." Interscience Publishers, New York, pp. 335-349.
- Cady, R. C., 1933, "Preliminary Report on Groundwater Resources of the Virginia Piedmont." Virginia Geological Survey Bulletin 41, 48 p.
- Cady, R. C., 1938, "Groundwater Resources of Northern Virginia." Virginia Geological Survey Bulletin 50, 200 p.
- Campbell, M. D., and Lehr, J. H., 1973, "Water Well Technology." New York, McGraw-Hill Book Company, 681 p.
- Darton, N. H., 1891, "Esozoic and Cenozoic Formations of Eastern Virginia and Maryland." Geological Society of America Bulletin, Vol. 2, pp. 431-450.
- Darton, N. H., and Keith, A., 1901, "Description of the Washington D.C.-Maryland-Virginia Quadrangles." U.S. Geological Survey Geological Atlas, Folio 70.
- Data Summary, Loudoun County, 1972, Department of Inter-governmental Affairs, Richmond, Virginia, 41 p.
- DeKay, R. H., 1976, "Hydrogeology" taken from, Geology of the Linden and Flint Hill Quadrangles, Virginia by M. T. Lukert and Ernest B. Nuckols, III. Virginia Division of Mineral Resources, Report of Investigation 44, pp. 55-64.

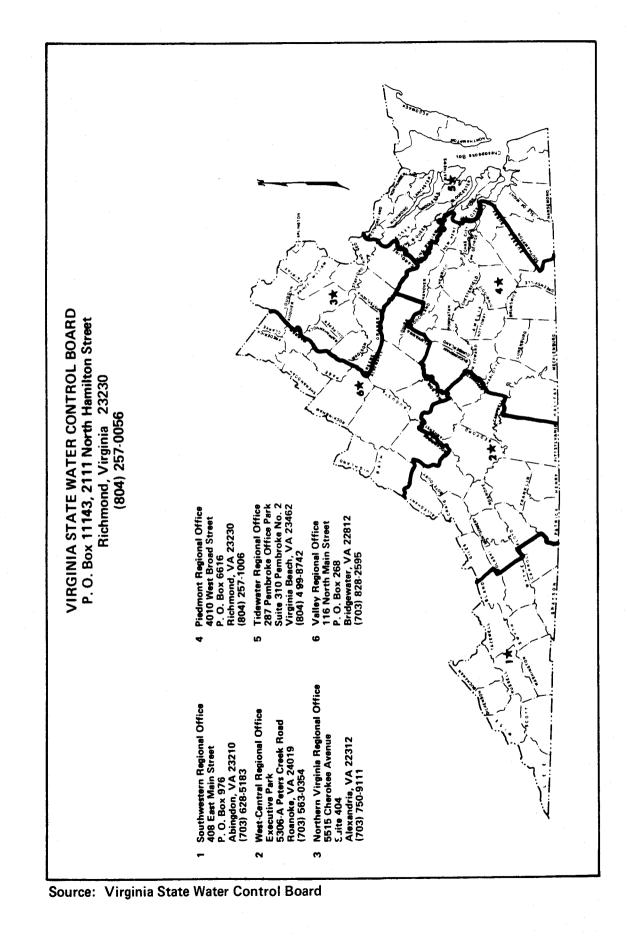
- Dewey, J. F., and Bird, J. M., 1970a, "Mountain Belts and the New Global Tectonics." Journal of Geophysical Research, Vol. 75, No. 14, pp. 2625-2647.
- Dewey, J. F., and Bird, J. M., 1970b, "Lithosphere Plate-Continental Margin Tectonics and the Evolution of the Appalachian Orogen, Geological Society of America Bulletin, Vol. 81, pp. 1031-1060.
- Dick, A. T., and Ball, L. B., 1945, "Some preliminary observations on the effect of Molybdenum on copper metabalism in herbivorous animals," Australia Verterinary Journal, 21:70.
- Dietrich, R. V., 1970, "Geology in Virginia." University of Virginia Press, Charlottesville, Virginia, 85 p.
- Environmental Studies Board, 1972, "Water Quality Criteria 1972." National Academy of Sciences, Washington, D.C., pp. 71, 73, 253, 314, 344.
- Espenshade, G. H., 1970, "Geology of the Northern Part of the Blue Ridge Anticlinorium, Studies of Appalachian Geology: Central and Southern." Interscience Publishers, New York, pp. 199-212.
- Fenneman, N. M., 1938, "Physiography of Eastern United States." New York, McGraw-Hill Book Company.
- Gass, T. E., 1978, "Drinking Water and Human Health: Part II." Water Well Journal, Vol. XXXII, No. 4, pp. 30-32.
- Gathright, T. M., and Nystrom, P. G., 1974, Geology of the Ashby Gap Quadrangle, Virginia." Virginia Division of Mineral Resources, Report of Investigations 36, 55 p.
- Geiger, H. R., and Keith, A., 1891, "The Structure of the Blue Ridge Near Harpers Ferry." Geological Society of America Bulletin, Vol. 2, pp. 155-164.
- Geraghty, J. J., and Miller, D. W., 1978, "Status of Groundwater Contamination in the USA." <u>Journal American Water Works Association</u>, Vol. 70, No. 3, pp. 162-167.
- Heiss, H. W., 1978, "Using large diameters to increase well storage." Water Well Journal, Vol. 32, No. 11, pp. 65&66.
- Hotchkiss, J., 1884, "The Copper Ores of Loudoun County, Virginia." The Virginians, Vol. 5, No. 12, p. 192.
- Howard, Ross., 1978, "Beware! This water can start a blaze." Toronto Star, November 25, 1978, pp. 1&16.

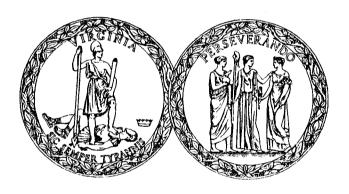
- Jonas, A. I., and Stose, G. W., 1939, "Age Relations of the Pre-Cambrian Rocks in the Catoctin Mountain - Blue Ridge and Mount Rogers Anticlinorium in Virginia." American Journal Science, Vol. 237, p. 573-593.
- Johnston, P.M., 1960, "Groundwater Supplies in Shale and Sandstone in Fairfax, Loudoun, and Prince William Counties, Virginia."
  U.S. Geological Survey Open File Report.
- Johnston, P.M., 1961, "Geology and Ground-water Resources of Washington, D.C. and Vicinity." U.S. Geological Survey Open File Report.
- Johnston, P. M., 1962a, "The Geology of Washington, D.C. and Vicinity." Washington Academy Science Journal, Vol. 52, No. 3, pp. 57-68.
- Johnston, P.M., 1962b, "Ground Water in the Washington, D.C. Area and a Brief History of the Public Water Supply." Washington Academy Science Journal, Vol. 52, pp. 101-110.
- Johnston, P.M., 1964, "Geology and Ground-Water Resources of Washington, D.C. and Vicinity." U.S. Geological Survey Water Supply Paper 1776, U.S. Government Printing Office, 97 p.
- Keith, A., 1893, "Geology of the Catoctin Belt." U.S. Geological Survey, 14th Annual Report, Part 2, pp. 285-395.
- Keith, A., 1894, "Description of the Harpers Ferry Quadrangle (Virginia, West Virginia, Maryland), U.S. Geological Survey Geological Atlas, Folio 10.
- Kellogg, C. E., 1941, "The Soils that Support Us." New York, 370 p.
- Lee, K. Y., 1977, "Triassic Stratigraphy in the Northern Part of the Culpeper Basin, Virginia and Maryland, Geological Survey Bulletin 1422-C, U.S. Government Printing Office, Washington, 17 p.
- Lee, K. Y., 1978, "Geologic Map of the Arcola Quadrangle, Loudoun, and Fairfax Counties, Virginia." Map MF-973, U.S. Geological Survey.
- LeGrand, H. E., 1954, "Geology and Groundwater in the Statesville Area, North Carolina." North Carolina Department of Conservation and Development, Division of Mineral Resources Bulletin No. 68, 68 p.
- Loudoun County League of Women Voters and the Chamber of Commerce, 1975, "Loudoun County, Virginia." 49 p.
- Mack, T., 1965, "Characteristics of the Everona Formation in Virginia." Virginia Division of Mineral Resources Information Circular 10, 16 p.

- McKee, J. E., and Wolf, H. W., eds., 1963, "Water Quality Criteria." 2nd ed., (California State Water Quality Control Board, Sacramento), pp. 132, 228.
- McKee, J. E., Laverty, F. B., and Hertel, R. M., 1972, "Gasoline in Groundwater." Journal Water Pollution Control, Vol. 44, No. 2.
- Meinzer, O. E., 1923a, "The Occurrence of Groundwater in the United States, Explanatory of a Geologic Map." American Philosophic Society, Trans. Vol. 5, Part 2, pp. 411-428.
- Meinzer, O. E., 1923b, "Outline of Groundwater Hydrology, with Definitions." U.S. Geological Survey Water Supply Paper 494.
- Miller, D. W., and Scalf, M. R., 1974, "New Priorities for Groundwater Quality Protection." <u>Groundwater</u>, Vol. 12, No. 6, November December.
- Parker, P. E., 1968, "Geologic Investigation of the Lincoln and Bluemont Quadrangles, Virginia." Virginia Division of Mineral Resources, Report of Investigations 14, 23 p.
- Reed, J. C., Jr., 1070, "Introduction, the Blue Ridge and the Reading Prong, Studies of Appalachian Geology, Central and Southern." Interscience Publishers, New York, pp. 195-197.
- Rogers, W. B., 1884, "Geology of the Virginias." New York, D. Appleton and Company, 832 p.
- Sowers, G. B., and Sowers, G. F., 1970, "Introductory Soil Mechanics and Foundations." The MacMillian Co., 556 p.
- Toewe, E. C., 1966, "Pyroclastic Rocsk in the Potomac Triassic Basin, Loudoun County, Virginia."
- Toewe, E. C., 1968, "Geology of the Leesburg Quadrangle, Virginia." Virginia Division of Mineral Resources, Report of Investigations 11, 52 p.
- U. S. Department of Housing and Urban Development, 1977, Federal Insurance Administration, "Flood Insurance Study, Loudoun County, Virginia, Unincorporated Area."
- U. S. Department of Agriculture, Soil Conservation Service, 1974, "Flood Hazard Analyses, Sycolin Creek, Loudoun County, Virginia." 12 p. plus photo maps 1-4, profiles 1-6.
- U. S. Department of Agriculture, Soil Conservation Service, 1974, Flood Hazard Administration, "Broad Run-Sugarland Run, Loudoun County, Virginia, 14 p. plus fig. 1, photo maps and flood profiles.

- U. S. Department of Agriculture, Soil Conservation Service, 1974, Flood Hazard Administration, "Tuscarora Creek, Loudoun County and Leesburg, Virginia." 14 p. plus fig. 1, photo maps 1-4, profiles 1-6, typical valley cross-section.
- U. S. Department of Agriculture, Soil Conservation Service, "Soil Survey, Loudoun County, Virginia." U.S. Government Printing Office, Washington, D.C., 1960.
- U. S. Environmental Protection Agency, 1976, "Quality Criteria for Water." Washington, D.C., 256 p.
- U. S. Geological Survey, 1976, "Water Resources Data for Virginia, Water Year 1975, U.S. Geological Survey Water-Data Report VA-75-1." U.S. Department of the Interior, Washington, D.C., pp. 46, 48, 49, 332.
- U. S. Geological Survey, 1977, "Water Resources Data for Virginia, Water Year 1976, U.S. Geological Survey Water-Data Report VA-76-1." U.S. Department of the Interior, Washington, D.C., pp. 42, 49, 50, 345.
- U. S. Geological Survey, 1978, "Water Resources Data for Virginia, Water Year 1977, U.S. Geological Survey Water-Data Report VA-77-1." U.S. Department of the Interior, Washington, D.C., pp. 46, 52, 53, 368.
- Virginia Department of Health, "Waterworks Regulations, Public Drinking Water Supply." Richmond, Virginia, May 1974.
- Virginia Division of Mineral Resources, 1963, "Geologic Map of Virginia." Charlottesville, Virginia.
- Virginia Division of State Planning and Community Affairs, "Population Projections, Virginia Counties and Cities, 1980-2000." 2000." Richmond, Virginia, March 1975.
- Virginia Division of Water Resources, 1969, "Information Bulletin 502." Department of Conservation and Economic Development, Richmond, Virginia, 86 p.
- Virginia Division of Water Resources, 1969, "Potoamc-Shenandoah River Basin Comprehensive Water Resources Plan." Volume I Introduction, Planning Bulletin 207. Richmond, Virginia.
- Virginia Devision of Water Resources, 1969, "Potomac-Shenandoah River Basin Comprehensive Water Resources Plan." Volume III -Hydrologic Analysis, Planning Bulletin 208. Richmond, Virginia.
- Virginia Division of Water Resources, 1969, "Potomac-Shenandoah River Basin Comprehensive Water Resources Plan." Volume IV Water Resources Requirements, Planning Bulletin 210. Richmond, Virginia.

- Virginia Division of Water Resources, 1969, "Potomac-Shenandoah River Basin Comprehensive Water Resources Plan." Volume V Engineering Development Alternatives, Planning Bulletin 211A, Richmond, Virginia.
- Virginia State Water Control Board, "The Groundwater Act of 1973 Chapter 3.4, Title 62.1, 1950 Code of Virginia, As Amended, Provisions for Administration, Rules, and Standards for Water Wells." 1973.
- Virginia State Water Control Board, 1973, "Groundwater in Virginia: Quality and Withdrawals." Bureau of Water Control Management Basic Data Bulletin 38, Richmond, Virginia.
- Virginia State Water Control Board, 1974, "Rules of the Board and Standards for Water Wells, 1950 Code of Virginia, As Amended." Richmond, Virginia, 33 p.
- Virginia State Water Control Board, 1974, "Guide for Water Well Contractors and Groundwater Uses." Bureau of Water Control Management Information Bulletin 508, Richmond, Virginia, 33 p. plus appendices A-H.
- Virginia State Water Control Board, 1976, "Water Quality Inventory, (305(b)) Report." Information Bulletin 52b, Richmond, Virginia pp. 67-112.
- Weber, R. S., 1975, "Soils of Loudoun County." Based on Soil Association Map of Loudoun County, by H. C. Porter, January 1952.
- Whitaker, J. C., 1955, "Geology of the Blue Ridge Near Harpers Ferry, West Virginia." Geological Society of America Bulletin Vol 67, No. 3, pp. 239-270.
- Williams, G. H., Darton, N. H., and Keith, A., 1901, "Description of the Washington, D.C.-Maryland-Virginia Quadrangle." U.S. Geological Survey Geological Atlas, Folio 70.





.